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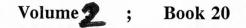
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# TOMBSTONE Mining District Cochise County, AZ

Santa Fe Mining, Inc. Semi-Annual Geologic Reports To T.D.C. 1988-1992 Excellon, U.S.A.



# SFM SANTA FE MINING INC.

# Albuquerque, New Mexico

## SEMI-ANNUAL REPORT

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TOMBSTONE DEVELOPMENT COMPANY

ON

EXPLORATION ACTIVITIES ON THE SFPM LEASE, TOMBSTONE ARIZONA

FEBRUARY 11 - AUGUST 11, 1988

# SEMI-ANNUAL REPORT TO TOMBSTONE DEVELOPMENT COMPANY ON EXPLORATION ACTIVITIES ON THE SFPM LEASE, TOMBSTONE ARIZONA FEBRUARY 11 - AUGUST 11, 1988

#### OVERVIEW

The purpose of the exploration program in the Tombstone area, to date, has been the gathering and integration of a variety of structural and geochemical data aimed at generating specific drill targets in the project area. The area of exploration, shown in Figure 1, was selected to evaluate numerous manifestations of possible carbonate replacement-type ore deposits such as those in the Charleston-Bronco Hill, State of Maine, and Section 32 areas, as well as in the main Tombstone mining district. The data gathering program has progressed through three major phases: a review of the literature and past exploration work, air photo and field structural interpretation, and geochemical sampling. This effort, to the date of this report, has resulted in the generation of two specific exploration targets, one of which is presently being drilled. Additional targets will be determined as the on-going process of data analysis continues. Thus far, approximately six man months of work has been expended on the project, at a cost of about \$52,000.00, of which approximately two thirds has been for work on the SFPM

UNITED STATES DEPARTMENT OF THE ARMY CORPS OF ENGINEERS

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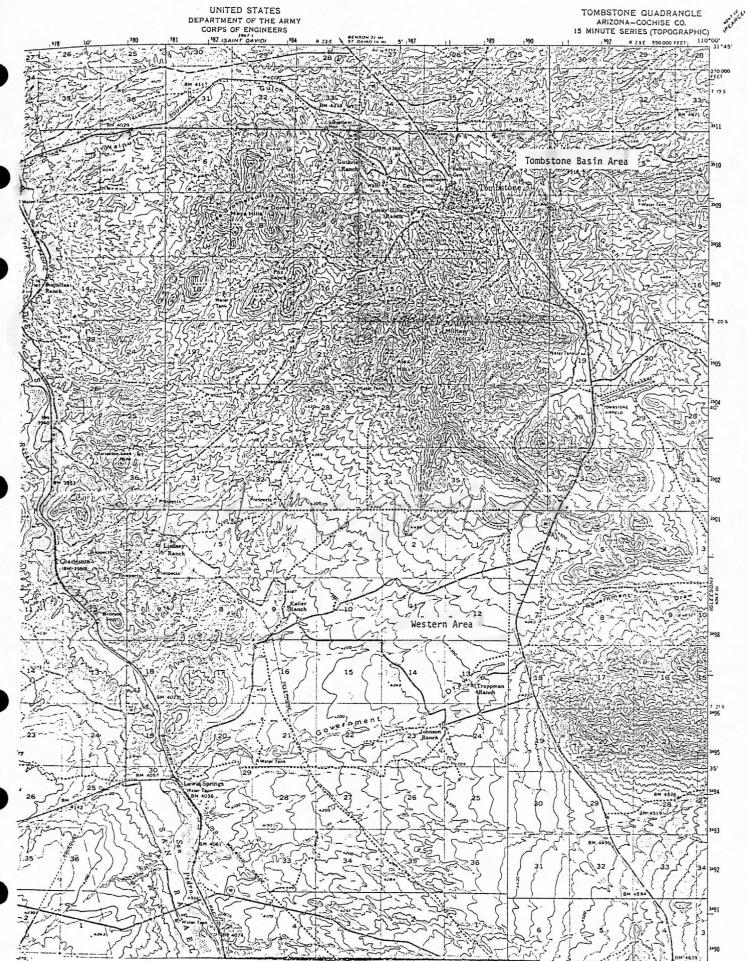


Figure 1. Proposed exploration areas in the Tombstone district and vicinity, Cochise County, Arizona.

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lease. The following sections outline the work accomplished in each of the exploration phases mentioned above, as well as discusses the drill targets generated thus far.

# LITERATURE SEARCH AND REVIEW OF PAST EXPLORATION

A thorough review of all of the published literature on the Tombstone district, as well as a review of available private archives, yielded valuable documentation of surface and sub-surface geology, ore controls, and certain geochemical zoning patterns. The information thus gathered strongly influenced the formulation and design of subsequent exploration program. Three major sources of information were available: published material, private company archives, and public collections of unpublished maps and reports.

The computerized GEOREF system was used to recover a comprehensive bibliography of the Tombstone literature dating from 1882 to the present. Included in the listing are early reports contained in various mining journals, state and federal articles and bulletins, articles in more recent geologic books, guidebooks, and journals, as well as theses and dissertations. Most valuable among this group is the dissertation by Newell (1975), which contains an excellent geologic map and abundant geochemical data, and the Arizona Bureau of Mines Bulletin 143 (Butler, <u>et</u>. <u>al</u>, 1938) which contains invaluable plans and cross sections of the ore bodies in the main Tombstone district.

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Little information was obtained from private company archives, with the exception of the map collection of Tombstone Development Company. Most companies who have mined or have done exploration work in the district either no longer exist, have destroyed or misplaced the pertinent information, or are unwilling to disclose what information they have. A case in point is Newmont Exploration whose personnel were unable to locate drill records or core from their 1952 underground program in the district. Interviews with retired Newmont engineers who performed the work also yield little useful information. The map files of Tombstone Development Company were the only source of private archival information which proved useful to the exploration program.

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Two archive collections, available to the public, contained the most valuable underground information obtainable from any source; those of the Arizona Bureau of Geology and the American Heritage Collection at the University of Wyoming. The former has a fragmentary collection of underground maps of various mines in the district, some showing underground geology, and a very few showing assay data. The underground workings and geology of the Lucky Cuss Mine are particularly well documented in this collection. The American Heritage Center's Anaconda collection, available to corporate users for an annual fee of \$8,500.00, contains an excellent collection of reports and mine maps covering the northeast part of the

Tombstone mining district. Level maps show detailed underground geology, details of mineralization, and abundant assay data.

The result of the literature search phase of the project was two-fold. First, it allowed the formation of several theoretical models for mineralization in the Tombstone district, and delineated numerous areas for field testing of these ideas. Secondly, the collected information, in part, dictated the design and methods to be used in the subsequent field program; i.e. it suggested the required areal extent of exploration program, as well as the type and density of field information required to identify specific drilling targets.

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## STRUCTURAL ANALYSIS

Insofar as many important carbonate replacement deposits are known to be controlled, at least in part, by structure, a thorough structural analysis of the Tombstone region was undertaken in order to identify the most promising areas for ground follow-up. Given the time and manpower available, it was decided that such analysis could best be undertaken using aerial photography. Various types of images were considered for use, ranging from LANDSAT and high-level U-2 images to low-level photography. Given the size and type of target under investigation, it was decided to utilize 1:24000 scale patural color aerial photographs.

The 1:24000 photos were obtained from Exploration Photography, Tucson, from an existing file. These photographs served well in delineating the major structural trends in the Tombstone area. Four sets of structural trends were documented: 1) an early, strong N-S trend primarily affecting Paleozoic and Cretaceous sedimentary rocks, 2) a strong E-W trend, again affecting Paleozoic and Cretaceous sediments, and to a lesser extent Cretaceous volcanics, 3) a NE trend strongly developed in Cretaceous sediments and volcanics, and 4) a late, weak NW trend affecting all rock types.

Despite attempts at computer enhancement, sufficient structural detail in areas of primary interest could not be obtained from the 1:24000 set which also suffered from poor color rendition. To rectify these problems, McLain Harbers Co., Inc. of Tucson, was contracted to re-fly the area to obtain a set of 1:12000 scale (from approximately 6000' above average terrain) natural color stereo photographs of the region, as well a three spot 1:24000 photographs from an altitude of 12000' above average terrain. The resulting set showed superb structural detail as well as excellent rock color rendition. Additional structural overlays were made on these photos and added to the previous 1:24000 interpre-Together, these were computer-transferred to 7.5 tations. minute topographic guadrangle maps. The 1:12000 photos were mapping additional subsequently used in the field for structural details and for plotting sample point locations.

The resulting structural interpretation map was then overlaid on a geologic map showing all mines, prospects, and mineralized trends in the Tombstone region. The resulting composite showed an excellent correspondence between certain structural trends and/or intersections, and areas of known mineralization. In addition, several areas of favorable structural intersection without <u>known</u> mineralization became apparent as the result of this analysis, serving to further focus the subsequent field mapping and sampling program.

# GEOCHEMICAL SAMPLING AND FIELD DATA COLLECTION

The bulk of the effort on the Tombstone project has been spent gathering geochemical samples, as well as collecting and compiling certain geologic data at the sample locations. A total of 355 surface and underground rock samples were collected in the project area previously defined by geologic and structural studies. Of these, 103 surface and underground samples were taken on lands leased by SFPM from Tombstone Development Company. Both surface and underground samples were also taken from property controlled by Tombstone Development adjacent to the SFPM lease; the results of this sampling will be included in the next semi-annual report of activities.

The goal of the sampling program was to obtain the geochemical signature of the highest grade mineralization in a given area, in order to generate metal ratio contour maps, raw assay contour maps, and/or metal value and metal ratio

vectors. To this end, samples consisted of either high-graded dump samples or channel samples of vein material in place. Considerable difficulty in obtaining such samples was experienced in the leased part of the main Tombstone mining district because of the extensive reworking of old dumps and general lack of underground access. An important exception to this situation was in the Lucky Cuss Mine area where sampling could be carried out underground on the 100-500 levels along a stretch between the Old Guard incline and the McCann patent.

All samples collected were assayed as follows: Au and Ag by standard fire assay (Copper State Labs, Tucson), 30 trace element suite by ICP analysis, As,Sb,Bi,Ge,Se, and Te by hydride generation ICP analysis, and F assay by NaOH fusion and ICP processing (Acme Analytical Laboratories, Vancouver). Assay results for samples collected on the SFPM lease are listed in Appendix I; sample locations are shown on the enclosed sample location map.

At each sample location, various geologic data were also collected. These include the following: type and dimensions of mineralization; sample mineralogy, paragenesis, and dilution; alteration type, mineralogy, width and intensity; host rock lithology; strike and dip of mineralized structure and cross structures; and estimated extent of mine workings. It is intended that all of this information will be encoded and entered into the computer data base, along with all assay data, for subsequent analysis. To date, however, only the assay data has been entered, as well as the coding for host

rock lithology, alteration type, and alteration intensity. These latter data appear in Appendix I, accompanying the assay results for each sample.

## DATA ANALYSIS

The goal of the data analysis phase of the project is the generation of specific drill target locations based on the integration and interpretation of all of the geologic, structural, and geochemical data at hand. Upon completion of data entry (which is still in progress), the following information will be available for each sample site: a three dimensional sample location (UTM coordinate and elevation), all assay data, and coded geologic/structural information, as discussed above. These data will be used to generate maps showing the distribution of raw assay data, contoured maps of various groups of metal ratios, and maps showing the various types of mineralization and alteration in the project area, and their relationships with each other. In addition, all assay data will be factor analyzed and sample scores plotted to show their areal distribution.

The end result of such analyses should be the delineation of structural/geochemical "bulls eye" targets and/or welldefined geochemical vectors pointing to mineralization centers both in the Tombstone main district and elsewhere in the area. Drill sites will then be selectively assigned to the highest quality targets.

#### DIAMOND DRILLING

Two drill hole sites have been established on the SFPM lease, the locations of which are shown on the enclosed sample location map. Vertical, 3000 foot deep, HQWL core holes are planned for both locations. Site locations were based on raw sample assay data, structural interpretation, and information gathered from the published literature.

Diamond drill hole T-1 was collared on July 22 at the common corner of the West Side, Blue Monday, and Tribute patents in an area of complex structure near the intersection of the West Side fissure and north-trending structures associated with the Tribute dike. The hole is designed to test for the presence of: 1) mantos, developed deep in the Naco section which are rooted in the West Side fissure feeder and/or 2) the development of a chimney along the West Side at a point of north-trending structural intersection(s). Secondary, the hole is intended to provide a good stratigraphic cross section of the middle and upper parts of the Naco group carbonates, and to test for the presence of Schieffelin granodiorite beneath this part of the mining district. A copy of the drill log for T-1 will be forwarded to Tombstone Development Company upon completion of the hole.

Drill hole T-2 is designed to explore the Lucky Cuss structure at a point approximately 2300 feet below the surface. To be collared about 700 feet due east of the Lucky Cuss shaft on the East Side patent, the hole is targeted on the intersection of the Lucky Cuss fault and the receptive

carbonates of the Middle Naco group, a potential site for the development of significant mantos or chimneys. The drill hole location was based on underground structural observations and the results of sampling in the Lucky Cuss Mine, as well as on published information which suggests an increase in gold values with depth in both the Lucky Cuss and Old Guard Mines.

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Additional drill holes are planned to test for deep mineralization on the Lucky Cuss structure. At present two holes are tentatively planned: one to be drilled beneath the Mn-Ag chimneys on the McCann and Wedge claims, and another beneath the Mn-Ag-Pb chimney cropping out at the Luck Sure Mine. Additional drill sites elsewhere on the SFPM lease will be spotted as data analysis progresses.

REPORT FEBRUARY, 1989

SEMI-ANNUAL REPORT TO TOMBSTONE DEVELOPMENT COMPANY ON ACTIVITIES ON THE SANTA FE PACIFIC MINING LEASE IN THE TOMBSTONE DISTRICT, COCHISE COUNTY, ARIZONA

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W. D. RIESMEYER FEBRUARY, 1989

# SEMI-ANNUAL REPORT TO TOMBSTONE DEVELOPMENT COMPANY ON ACTIVITIES ON THE SANTA FE PACIFIC MINING LEASE IN THE TOMBSTONE DISTRICT, COCHISE COUNTY, ARIZONA

# W. D. RIESMEYER

#### SUMMARY

A total of five core holes, with an aggregate footage of 14,841 feet, was drilled on the Santa Fe Pacific Mining lease in the Tombstone district between July 21 and December 20, Several mineralized intercepts were obtained but none 1988. were of both ore grade and width. In addition to the Cretaceous Bisbee formation, an extensive Paleozoic stratigraphy was cored during the program. This included the Martin, Escabrosa, Horquilla, Earp, Colina, and Epitaph While all formations contained mineralization, formations. drill results showed the Earp to be most strongly mineralized, followed by the Epitaph and Colina. Carbonates, particularly recrystallized and/or skarned limestone, appear to be the most receptive host for replacement type mineralization which occurs only rarely in the other lithologies cored during the program. Structural observations suggest that formations occurring deep in the section and consisting primarily of carbonates (i.e. the Martin through the Colina) deformed plastically, while those higher in the section, and having a large clastic component (i.e. the Epitaph the Bisbee), deformed in a more brittle fashion. Mineralization occurring in the core may be classified into six types: 1) base metal

replacements, 2) pyrite replacements, 3) skarn-related mineralization, 4) fracture-controlled mineralization, 5) manganese oxide replacements, and 6) base metal oxide deposits. Of these, types 1, 3, 5, and 6 are volumetrically most important and have yielded the best assay results. Types 1, 2, 5 and 6 resemble manto-type replacements. Some are high are too thin to constitute ore. Type 4 grade, but mineralization includes guartz vein deposits which are of high grade but are volumetrically insignificant. Type 3, skarn mineralization, while widespread is generally of very low grade. Very little direct structural information was derived from the drilling program aside from confirming the continuity of the Lucky Cuss fault to depths exceeding 2,500 feet. Two drill holes missed their targets, apparently because of significant changes in the dip, at depth, of major structural features such as the Westside fissure and the Empire dike. Inter-hole stratigraphic correlations suggest the presence of major, pre-Bishee faulting which may be significant in planning further exploration in the district.

#### INTRODUCTION

This report outlines and discusses the results of the drilling program conducted by Santa Fe Pacific Mining, Inc. between July and December, 1988, on lands leased from Tombstone Development Company. The drilling program represents a part of a continuing exploration effort in and around the Tombstone mining district which began in early 1988 with an extensive mapping and sampling project, the results of which were outlined in a previous semi-annual report of The recently completed phase of activities. drilling consisted of five deep core holes with an aggregate footage of 14,841, targeted on areas thought to be favorable for the occurrence of economic carbonate replacement deposits. While numerous mineralized intervals were discovered in the course of the program, none were found to be of adequate thickness and/or grade to be economic. The drilling did, however, yield information pertaining to stratigraphy and structure which will be of value in targeting additional core holes on the property.

This report is intended as a supplement to the accompanying detailed drill logs for holes T1-T5 which record the stratigraphic and structural information from each hole, as well as descriptions and assay data for the mineralized intervals encountered. Laboratory assay sheets are appended. The following report expands on the log information. The first section describes, in detail, the stratigraphy revealed

by the drilling program. Included are lithologic descriptions and paleontological information, as well as summaries of the alteration, structure, and mineralization found in each formation. The second section classifies and describes the various types of mineralization found in the drill core. Each type is described in terms of form, mineralogy, related rock. alteration, structure, preferred host and Representative metal values for each type of mineralization are also included. The final section discusses the structure observed in the core and also describes the structural geology implied by hole to hole correlations. The report is accompanied by two plates. Plate 1 is a geologic map of the Tombstone district which shows the results of mapping during the first phase of exploration, as well as the location of the five drill holes. Plate 2 presents simplified graphic logs for each hole. Summary structural and mineralization data are also plotted on the logs, and inferred inter-hole correlations are shown.

#### PROGRAM OUTLINE

A total of 14,841 feet of core was drilled in the Tombstone district between July 21st and December 20th, 1988. Five holes, designated T1-T5, were completed. DDH T1 was drilled on the East Side (280) patent to a depth of 2,834. The hole was lost short of its projected TD at 3000 feet because of caving ground. T2, drilled on the Elue Monday claim, reached TD at 3,003 feet. T3, on the East Side (376), was TDed at 3,000 feet. T4 was drilled on the Luck Sure and reached a depth of 3,000 feet. T5 was collared on the Silver Thread patent and TDed at 3,004 feet beneath the Empire claim. All holes were vertical except T5 which was drilled N57<sup>o</sup>W at -80<sup>o</sup>. NC core was recovered from the upper parts of the holes until reduction became necessary to NX core. Core recovery in all holes averaged 100%.

Drilling was carried out by Longyear Contract Drilling Company utilizing Longyear Hydro-44 rigs. The program began using two 10 hour shifts. Inadequate progress, however, required changing to three 8 hour shifts in early November. Overall average production was approximately 35 feet per shift. Core logging was done on site on a daily basis. W. D. Riesmeyer and K. E. Eagan (North American Exploration, Inc.) where the primary core loggers. P.K.M. Megaw logged the majority of T2 early in the program. The core was stored on site in a watertight sea-going container. Core for assay was shipped to various labs for splitting or sawing, and sample prep.

Assaying, early in the program, was done by Copper State Analytical Laboratory in Tucson, but was later shifted to Mountain States R&D in Vail, Arizona. Gold and silver were determined by standard fire assay; copper, lead, and zinc by regular assay. Thirty trace elements were determined by ICP analysis, and As, Sb, Bi, Ge, Se, and Te by hydride generation and ICP analysis. Fluorine was determined by sample fusion with NaOH, leaching and analysis by specific ion electrode. All minor element determinations were done by Acme Analytical Laboratories, Ltd., Vancouver, British Columbia. Assay results were entered in the computerized Tombstone data base, and major metal assay values were posted on the drill logs.

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### STRATIGRAPHY

The following formations were cored during the 1988 drilling program: Martin limestone, Escabrosa limestone, Horquilla limestone, Earp formation, Colina limestone, Epitaph dolomite, and the Bisbee formation. In addition, Schieffelin granodiorite was encountered in one hole. In general, the Bisbee and formations of the upper Naco group were recovered in drill holes 1-4, while the Bisbee, Horquilla, Escabrosa, and Martin were cored in hole 5 in an apparently up-faulted block.

Most rocks in the project area are moderately to strongly altered. This, along with the paucity of fossils, made the precise determination of formation tops and bottoms difficult. Formation contacts could, however, be determined with a fair degree of certainty within a few to a few tens of feet. The stratigraphy of each hole, as well as hole to hole correlations, are summarized on Plate 2.

#### Martin Limestone

Distribution and Thickness. The upper Devonian Martin limestone was encountered only in DDH T5 where it was cored in the lower 339 feet of the hole. Only two questionable bedding angles were determined which suggest between 185 and 300 feet true thickness of the Martin were cored. Gilluly measured a true thickness of 230 feet in the Tombstone Hills, suggesting a nearly complete section may have been drilled in T5.

Lithology. In drill hole T5, the Martin consists of light to medium gray, silty to argillaceous, generally massively bedded limestone, with abundant interbedded siltstone. The siltstones form beds from a few inches up to 30 feet thick consisting of tan, brown, or green, fine to medium grained clastic material. Such beds are far more abundant than in the overlying Escabrosa. The limestone generally lacks any distinguishing characteristics except for the presence of black (organic) stylolites which are common in the lower part of the unit.

Stratigraphic Relationships. The Martin is in apparent conformable contact with the overlying Escabrosa. The upper contact was placed at the contact between "clean" Escabrosa limestone and the distinctly silty limestone of the Martin. The abundance of interbedded siltstone also served to differentiate the two formations. Scattered intervals of very poorly preserved crinoid stems and brachiopods are present in the cored section, but they were of no diagnostic value.

Alteration Summary. Due to the silty and argillaceous nature of the limestone, skarn development is prominent in the Martin. Several long intervals of weak to moderate wollastonite garnet-vesuvianite skarn with a pseudoconglomerate structure were cored. Where not skarned, the limestone is moderately recrystallized. Siltstone units are moderately hornfelsed and contain significant amounts of epidote. ?

Mineralization Summary. No significant mineralization was found in the Martin. Weak, disseminated and veinletcontrolled pyrite mineralization is associated with skarned intervals, as are trace amounts of sphalerite. Two rhodochrosite "veins" (2" and 6"), with associated trace pyrite and sphalerite, were cut in the lower part of the formation.

Structural Characteristics Summary. In T5, the Martin is "tight"; fracture intensity is low, both in limestone and siltstone units, even where skarned or hornfelsed. Fracture angles are generally  $>45^{\circ}$  to the core axis. The only exception to these characteristics is a 15 foot sheared and brecciated zone in an unaltered siltstone unit high in the formation. Here, nearly vertical fractures and shears are tightly cemented by calcite.

# Escabrosa Limestone

Distribution and Thickness. The lower Mississippian Escabrosa limestone was cored only in DDH T5 where a total of 702 feet was recovered. Measured bedding angles suggest a true thickness of 450 feet was cut. This does not match well with Gilluly's measurement of 733 feet in the Tombstone Hills, perhaps indicating a part of the T5 section has been faulted out or attenuated along the limb of a fold.

Lithology. The Escabrosa in T5 is primarily a limestone unit, but, unlike the formation elsewhere in Cochise County, it here contains approximately 12% siltstone interbeds. Where

least altered, the limestone is very light to medium gray, fine grained, and massively bedded; some intervals consisting of crinoid cocina. Several sections contain one to two inch, round to irregularly shaped, dark gray to cream colored chert nodules. Chert is far less common, however, than in the overlying Horquilla. Limestones are weakly stylolitic towards the top of the formation. Siltstone interbeds range from a few inches up to 13 feet thick. They are predominantly brown, but are also black, green, or dark gray, and are fine to medium grained, and thinly bedded to massive. One five foot bed of fine grained, greenish white dolomite was cored near the middle of the formation.

Escabrosa is in Stratigraphic Relationships. The apparent conformable contact with the underlying Martin and the overlying Horguilla limestone. The upper contact is dramatic increase in the degree of distinguished by a recrystallization in Escabrosa limestone, compared with that in the Horguilla, as well as a substantial decrease in the While this amount of chert downward across the contact. division may appear to be guite arbitrary, the contact, as picked, fits very well with published descriptions. The fossil content of the Escabrosa is rather sparse, consisting largely of crinoid stems and unidentifiable brachiopods, gastropods (up to 1" tall), rugosa (up to 3" long), and various pieces of trilobites. One cross section of a pygidium, sufficiently well preserved and distinctive, allowed

positive identification of the lower Mississippian trilobite crytosymbole hercules, confirming the unit to be Escabrosa.

Alteration Summary. The entire Escabrosa section in T5 is altered to one degree or another. The upper half of the unit is strongly recrystallized, yielding a coarse grained white marble. Interbedded siltstones are strongly hornfelsed. In the lower half, limestones are either weakly recrystallized or weakly to moderately skarned. Where best developed (around and in cherty horizons), the skarn has a pseudo-conglomerate structure with mottled pink, apple green, and gray colors, and silica, tan garnet, epidote, chlorite, and consists of wollastonite. Other intervals are silicified in a distinctive anastomosing network fashion. Also, in the lower half of the formation, short two and three foot intervals have been converted to light green talc probably derived from dolomite Siltstone units in the lower half are weakly to interbeds. moderately hornfelsed; some hornfels intervals have been bleached.

<u>Mineralization Summary</u>. Disseminated and veinletcontrolled pyrite occurs sparingly throughout the Escabrosa, most commonly associated with hornfels and the most strongly developed skarn intervals. Similarly, very weak sphalerite, galena, and/or chalcopyrite disseminations or veinlets occur, associated usually with skarn intervals. Nowhere do base metal sulfides exceed 1%. Two 2" rhodochrosite "veins" occur in the cored section which are accompanied by trace sphalerite, galena, and chalcopyrite.

Structural Characteristics Summary. The Escabrosa, in T5, is very "tight"; fracture intensity is very low throughout, except in the lone dolomite bed where it is moderate. Most fracture orientations fall between  $30^{\circ}$  and  $60^{\circ}$ to the core axis. Faults, shears, breccias, and other signs of brittle deformation are altogether lacking, while evidence of plastic flow, with nearly vertical orientation, is abundant.

## <u>Naco Group</u> Horguilla Limestone

Distribution and Thickness. As is the case with the above-mentioned units, the lower Pennsylvanian Horquilla limestone was cored only in DDH T5. A total of 1,497 feet was cored. Measured, well-defined bedding shows that T5 cored a true thickness of 1,060 feet which compares very well with the 999 foot thickness measured by Gilluly in the Tombstone Hills.

The Horquilla in T5 consists of over 95% Lithology. The remainder consists of siltstone/shale beds up limestone. cherty foot bed of and one four feet thick, to 16 (intraformational ?) conglomerate. bulk of the The siltstone/shale, as well as the conglomerate, occur in the limestone, where upper guarter of the formation. The unaltered, is very light gray, fine to medium grained, and massive to thickly bedded. Several intervals in the upper guarter have a distinct shaly component, while the lower three quarters of the formation is very "clean" limestone. Large, brown to black, rounded chert nodules occur sparingly in the upper quarter of the Horquilla, are rare in the middle parts, and are guite numerous in the lower guarter. The general abundance of chert in the Horquilla serves well to distinguish it from limestones of the overlying formations of the Naco Group.

Stratigraphic Relationships. The upper contact of the Horquilla in T5 is a profound erosional unconformity being overlain by the Cretaceous Bisbee formation. While the lowermost Bisbee is a conglomerate, there is no evidence, in the core, of an erosional surface at the top of the Horguilla. The definition of the lower Horguilla contact has been discussed previously. In addition to the abundant chert. identification of the Horquilla was materially aided by the presence of abundant fossils. The formation contains numerous large, chert filled burrows and other evidence of extreme bioturbation. Also observed were long sections composed "trash": almost exclusively of fossil crinoid stems. fusilinids, rugosa, brachiopod fragments (both articulated and non-articulated types), and bryozoa. One rugosa was identified with fair certainty as rugosa lophophyllidium a Penn-Perm horn coral. Most important, however, was the discovery of the lower Pennsylvanian index fusilinid triticites sp? exceptionally well preserved in chert in the lower part of the formation.

Alteration Summary. Alteration of the Horquilla in T5 consists primarily of recrystallization and skarning of the limestone and development of hornfels in the rare siltstone units. Recrystallization affects the upper and middle parts of the formation. In these areas, long sections have been converted to a coarse grained gray-white marble. In the lower half, only patchy, weak recrystallization was observed. Stashows similar spatial development. Unlike most other

formations in the project area, Horquilla skarn mineralogy is simple, consisting of quartz and wollastonite, with only rare, patchy development of garnet.

Mineralization Summary. Most mineralization observed in the Horduilla occurs in the upper half of the formation associated with skarn or recrystallized limestone. Here, several zones of manganese oxide dendrites with associated limestone alteration and disseminated purite were pink Zones of trace to very weak, disseminated and observed. veinlet-controlled pyrite, sphalerite, and galena in skarn are also common in the upper half of the Horquilla, but none returned significant assay values. The best intercept was a ten foot segment of galena>sphalerite>>pyrite associated with a guartz vein in recrystallized limestone, four feet of which ran 8.99 opt Ag and .06 opt Au.

Structural Characteristics Summary. Like the carbonate formations below it, the Horquilla is "tight", with a low fracture density throughout. Fracture angles are low; generally less than  $45^{\circ}$  from horizontal; high angle fractures are exceedingly rare. What best exemplifies the structural character of the Horquilla is the presence of numerous zones of plastic deformation and flowage, all of which are oriented within  $15^{\circ}$  of vertical. This strongly suggests that the Horquilla deformed plastically along the limbs of the Silver Thread anticline rather than yielding in a brittle fashion.

#### Earp Formation

Distribution and Thickness. The upper Pennsylvanian Earp formation was cored in DDHs T1, T2, T3, and T4. All four holes were TDed in Earp sediments. Footages cored were 225, 455, 585, and 1,250 feet respectively. Measured bedding angles indicate true thicknesses of 180, 410, 480, and 715 feet. Gilluly reports a thickness of 595 feet for the Earp in the Tombstone Hills. The great thickness of the formation in T4 is undoubtedly due to repetition of the unit by faulting, of which there is abundant evidence in the core.

Lithology. The Earp is easily distinguished from other members of the Naco group by its large clastic component. This component ranges from 74% of the formation in T2 to 34% in T4, and consists mostly of shale and siltstone, along with one short interval and sandstone and one of arkose. Earo shales are black to greenish black, laminated to thinly bedded, and generally calcareous. Siltstones range in color from tans and browns to grays and blacks. Most are finely bedded, and many intervals are calcareous. The lone sandstone unit, in T3, is light gray, very fine grained, and thinly There is no apparent pattern of upward fining or bedded. coarsening of the clastic units within the Earp. Earp limestone units are light shades of gray where unaltered, and white where recrystallized. Unaltered limestones are fine to medium grained and generally massive, although rare bedding planes are occasionally observed. Much of the limestone in the Earp section is argillaceous or has a silty component,

judging by the abundant development of skarn minerals. Stylolites are common in several limestone intervals. A few dolomite beds, up to 10 feet thick, are interbedded with limestone in two holes.

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Stratigraphic Relationships. The top of the Earp is placed at the top of the first significant clastic unit occurring beneath the Colina. While an occasional similar clastic bed occurs within the Colina, the character of any underlying Colina limestone is so distinctive that picking the first Earo clastic unit can be done with confidence. In core, the Colina-Earp contact is entirely conformable; there was probably no time break between the deposition of the two The bottom contact, that between the Earp and formations. Horquilla, was not penetrated in any drill hole. The Earp is only sparingly fossiliferous and those fossils that exist are very poorly preserved. Crinoid stem and rugosa fragments were found, along with unidentifiable fusilinids; none were of any use in determining the age of the formation.

<u>Alteration Summary</u>. Irregular development of recrystallization, skarn, and hornfels characterize the alteration of the Earp. Alteration patterns are complex within a single hole and are not consistent between holes, undoubtedly due to the complex vertical and lateral changes between carbonate and clastic lithologies. The Earp in T1 shows moderate to strong recrystallization and hornfels development, but no significant skarn, while T4 shows very strong recrystallization and hornfelsing and patchy skarning. T2 and T3 show similar

alteration, but also contain intervals of skarn which appear to increase in thickness and intensity down hole. Skarn in the lower Earp of T2 and T3 is the most strongly developed and the most mineralogically complex of any found to date in the Developed both in limestones and in calcareous district. clastics, skarning is so intense that little or no remnant carbonate remains in the core. In these areas, the rock has been totally converted to guartz, green and/or tan garnet, In some skarn zones a brown vesuvianite, and wollastonite. In addition to these prograde pyroxene is also present. minerals, retrograde (?) actinolite, chlorite, talc, and epidote are present in some skarn zones.

Mineralization Summary. The Earp is perhaps the most heavily mineralized formation drilled, although nowhere were ore grade intercepts obtained. Pyrite mineralization occurs with all lithologies and alteration types, but shows a strong and/or hornfelsed siltstone. preference for calcareous Disseminated to semi-massive manto-type galena-sphalerite+ chalcopyrite mineralization mainly in skarned occurs limestone; one 1 foot interval of such mineralization assayed 13.42 opt Ag, 5.10 Pb, and 3.66 Zn. Similar mineralization also occurs in skarned dolomite, and, to a leaser degree in limestone. Perhaps the most intriguing recrystallized mineralized interval occurs in probable cave fill below a 13 foot void in T4. Here, at a depth of 2845, feet wulfenite, mimetite, hemimorphite, and possible bromargyrite are found

disseminated through the vuggy fill. Manganese oxides appear to form a halo above and below the mineralization.

Structural Characteristics Summary. Overall, the fracture intensity in the Earp is low. However, unlike lower formations, there exist several zones of moderate to high fracture density, especially in the clastic units and in the occasional dolomite horizon. Fracture orientation spans the range from nearly flat to vertical. Further, the Earp is broken by numerous faults which have left relatively open gouge and breccia zones. No evidence of plastic deformation was observed. In general, the Earp appears to have yielded in a more brittle fashion than underlying units.

## Colina Limestone

Distribution and Thickness. The Permian Colina limestone was cored in DDH's T1, T2, T3, and T4. In these holes, a total of 675, 655, 680, and 510 feet of the formation, respectively, was recovered. Recalculation, using very questionable bedding angles, shows 520, 505, 490, and 335 feet true thickness were cored. Gilluly reports a thickness of 635 feet in the Tombstone Hills.

Lithology. The Colina is the most distinctive formation cored during the drilling program. With few exceptions, it is a white to very light gray, coarse grained calcite marble which, in places, is stylolitic. Only in a few short intervals does recrystallization decrease enough to hint of the Colina's original character - a dense, aphanitic, black

limestone. In DDH T2 and T4, the Colina consists exclusively of limestone marble. In T1 and T3, a few beds of thinly bedded dark gray siltstones and shales, generally less than 10 feet thick, occur towards the bottom of the formation, reflecting a change towards the more clastic sedimentation of the underlying Earp.

<u>Stratigraphic Relationships</u>. The top of the Colina is placed at the base of the last clastic or dolomite unit of the overlying Epitaph formation. Because of the distinctive marblized nature of the Colina and the lack of clastic units in its upper part, the top contact is easily spotted. The character of the lower contact with the Earp has been previously discussed. Because of the strong recrystallization, most fossils in the Colina have been destroyed. None of the Colina-characteristic gastropods <u>omphalotrochus</u> were observed in the core; the only fossils found were poorly preserved and unidentifiable fusilinids.

<u>Alteration Summary</u>. Strong recrystallization is the most characteristic alteration type in the Colina and has affected all of the limestone to one degree or another. Siltstone/shale horizons are strongly hornfelsed. Towards the bottom of the unit, numerous zones of generally weak guartzwollastonite-garnet-vesuvianite<u>+pyroxene+actinolite+epidote</u><u>+</u> chlorite skarn have been developed in what must originally have been zones of argillaceous limestone. Short intervals (<5') of Colina marble have been sanded, yielding a mass of

extremely friable calcite crystals. Significant sanding has not been observed in any unit other than the Colina.

Mineralization Summary. The Colina is the third most well mineralized unit next to the Earp and Epitaph. Numerous thin mantos or thin, bedding-parallel veins of massive galenasphalerite-chalcopyrite+pyrite are scattered throughout the They range from paper thin up to a few inches formation. thick. One such "mini-manto", 1.5 inches thick, assayed 1.87 opt Ag, 2.14% Cu, 10.4% Pb, and 5.4% Zn. Finely disseminated galena-sphalerite-pyrite occurs associated with skarned intervals in the lower part of the formation, and along. organic-rich stylolites. In either occurrence sulfides do not Significant manganese oxide replacements occur in exceed 1%. the Colina in T1 and T4. In T1, a seven foot section is largely replaced by manganese oxides and only assays .35 ppm In T4 there are numerous, thin manganese oxide mantos. Ag. Two such zones, two and six feet thick, contain, in addition to manganese minerals, wulfenite, mimetite, and fluorite; assemblages similar to those found in stopes of the Empire Mine. One 1 foot section assayed 2.84 opt Ag.

Structural Characteristics Summary. Fracture density in the Colina is generally low and fracture orientation is usually  $>45^{\circ}$ . The intensity of fracturing tends to increase in skarned intervals and in siltstone interbeds. There are few faults in the Colina; those which exist have left tight, thoroughly healed breccias or shears. Any convincing evidence

of pre- or syn-metamorphic plastic or brittle deformation of the Colina has been obliterated by the intense recrystallization.

## Epitaph Dolomite

Distribution and Thickness. The Permian Epitaph dolomite occurs in drill holes T1-T4. The cored footages are as follows: T1: 1,010', T2: 540', T3: 1,040', and T4: 590'. Adjusted for bedding angles, true thicknesses are 780', 440', 735', and 200' respectively. Gilluly reports a thickness of 780 feet in the Tombstone Hills. The abnormal thinness of the formation in T4 is undoubtedly due to intense faulting along the Lucky Cuss structure. The thin section in T2 may be more apparent than real in that the upper contact with the problematical Scherrer formation may be misplaced, or part of the Epitaph may simply have been faulted out (see Plate 2).

Lithology. Epitaph lithologies are extremely varied. Included are limestone, dolomite, shale, siltstone, arkose, sandstone and conglomerate. The reader is referred to Plate 2 for a generalized view of the distribution of the various lithologies in the drill holes. It would appear that facies changes in the Epitaph were quite rapid. Thick sequences of dolomite, for example, are present in T1 and T2, especially in the upper and lower quarters. Dolomite is lacking, however in T3 and T4, suggesting either a rapid facies change or localized dolomitization. Clastic units are abundant in T1, T3, and T4, but are almost absent in T2, while carbonates are

abundant in all holes except T4 which is composed to a large degree of siltstone and arkose.

Stratigraphic Relationships. The upper contact of the Epitaph rests against basal conglomerate of either the Bisbee Scherrer formations clearly an erosional and is or The Bisbee basal conglomerate is composed unconformity. primarily of limestone fragments mixed with rarer sandstone and siltstone pebbles, in an arkosic matrix. The Scherrer basal conglomerate consists of epidotized dolomite clasts. The nature of the lower contact of the Epitaph has been discussed above. Only a very few fossils were found in the formation: forams, fusilinids, brachiopods, and echinoderm spines; none were of diagnostic value.

Alteration Summary. In general, the Epitaph shows a noticeably lower degree of pervasive alteration than other formations in the project area. While certain units in the Epitaph show very strong skarn or hornfels development, the overall aspect of alteration in the formation is weak. The exceptions to this observation are a strong skarn section in T3, and moderate to strong argillization found in T2 and T4. The skarn section in T3 is unique among those in other drill holes because of its thickness - slightly more than 200 feet, and its distinctive structure - that of a pseudo-conglomerate. Skarning apparently progressed along bedding and fractures to yield a rock with rounded "clasts" set in a matrix of skarn, all of which resembles a sedimentary conglomerate. Skarn minerals are guartz, wollastonite, tan garnet, and

vesuvianite, along with rare brown pyroxene. In DDH T-2, Epitaph limestone has been largely converted to a white to tan gritty clay for several feet above and below a mineralized interval. Save for a short section of equally intense, but non-mineral-related argillization in T4, this type of alteration is also unique among all the drill holes.

Mineralization Summary. The Epitaph is the second most intensely mineralized formation after the Earp. Unlike the Earp, mineralization in the Epitaph tends to occur in discrete thin mantos and is not scattered throughout the formation as veinlets and disseminations. In T1, a semi-massive manto, two feet thick, composed of galena and sphalerite, assayed 5.81 opt Ag, .09% Cu, 8.3% Pb, and 7.5% Zn over the best one foot. A similar two foot thick manto in T2, which contained, in addition, significant chalcopyrite, assayed 1.32 opt Ag, 1.04% Cu, 5% Pb, and 3.95 Zn over one foot. Several semi-massive pyrite mantos, one to two feet thick, also occur in T1, but all failed to return significant assays. T3 and T4 lacked is significant that a11 noteworthy mineralization. It important mineralization observed in the Epitaph occurs in limestone units and noticeably avoids adjacent dolomite beds.

<u>Structural Characteristics Summary</u>. The Epitaph has yielded in a far more brittle manner than any of the underlying formations. Fracture intensity, while highly variable, is strong over many intervals, and shows a variety of orientations. Near vertical orientations are more common in the Epitaph than in underlying units. Numerous faults and

shears are present in the formation, many of which have left partially open breccias and semi-porous gouge zones. Only in the thickest carbonate units is fracture intensity low, but even here, there is no strong evidence of plastic deformation.

## Scherrer Formation (?)

Distribution and Thickness. The Permian Scherrer formation (?) was observed in only one drill hole: T2. Its presence and identification are somewhat problematic in that Scherrer has never been reported in the Tombstone the district; its closest documented occurrence is at the northern end of the Dragoon Range. A total of 895 feet of probable Scherrer was cored in T2 with an estimated true thickness of 770 feet. Gilluly reports a thickness of 687 feet in the Gunnison Hills.

Lithology. The Scherrer is primarily a clastic formation with limestones, dolomitic limestones, and dolomites occurring in the middle and upper parts. The lower half of the formation consists of a generally coarsening upward sequence of lower shales and upper siltstones. The upper half of the formation consists of sandstones with thick interbeds of limestone or dolomite. In the sandstone units, several thick graded sequences were observed consisting of basal pebble conglomerate ranging upward to silty, fine grained sandstones.

Stratigraphic Relationships. The top of the Scherrer in T2 is well-defined at the base of the basal (Glance) limestone conglomerate of the Bisbee formation. The contact is clearly

erosional, but there is no apparent angular unconformity between the two formations. The lower contact of the Scherrer is set at the base of a dolomite pebble conglomerate which overlies shales of the Epitaph formation. The placing of this contact is somewhat arbitrary and is perhaps too low, accounting for the anomalously thin section of underlying Epitaph referred to previously. No fossils were found in the Scherrer to assist in resolving this problem.

Alteration Summary. Alteration throughout the Scherrer is generally weak with the exception of moderate recrystallization of some of the carbonate units. Clastic units, especially those deepest in the section, show weak development of hornfels. Weak wollastonite-actinolite skarn is patchily developed in some carbonate units and in sections of calcareous sandstone. With the exception of the Bisbee, the Scherrer is the least altered formation observed during the present drilling program.

<u>Mineralization Summary</u>. Weak disseminated and veinletcontrolled pyrite mineralization is found throughout the Scherrer. Weak manganese oxide replacements of carbonate units are scattered through the upper half of the formation. No significant base or precious metal mineralization, however, occurs anywhere within the Scherrer.

<u>Structural Characteristics Summary</u>. Brittle deformation characterizes the Scherrer. Fracture density, while variable, is high over long sections. Fracture orientation is also highly variable with numerous high and low angle sets

observed. Numerous breccias, both semi-open and healed are present, most notably in the brittle sandstone units. Carbonate units, while showing low to moderate fracture density, host numerous breccias and shears, and show no evidence of flowage or plastic deformation.

#### Bisbee Formation

Distribution and Thickness. The Cretaceous Bisbee formation overlies Paleozoic rocks in all drill holes in the project area. Cored thicknesses are as follows: T1: 650', T2: 445', T3: 685', T4: 650', and T5: 395'. Calculated true thicknesses are 330', 295', 310', 120', and 250' respectively. In the Tombstone area, Gilluly reports a thickness is excess of 3000 feet. Accordingly, all drill holes penetrated the very basal parts of the Bisbee.

Lithology. In the project area, the Bisbee formation consists of a basal conglomerate (Glance) of variable thickness overlain by finer grained clastic rocks with The basal conglomerate ranges from 8 interbedded limestones. to 28 feet thick. In T2, T3, and T4, limestone pebbles and cobbles are the dominant clast lithology, while sandstone and siltstone fragments are the primary clast types in T1 and T5. Overlying the conglomerate are thick, monotonous intervals of siltstone or arkose with lesser amounts of sandstone and Graded bedding is common in some units. conglomerate. Limestone units are scattered throughout the section, but are more numerous towards the bottom. They range in thickness from a few inches up to 75 feet. Hole to hole correlation of

the limestone units proved impossible (see Plate 2), suggesting the carbonate units are very lenticular and that the historically applied names (eg. the "Blue", "Joe", etc.) are of use only on a very local basis.

<u>Stratigraphic Relationships</u>. As discussed above, the base of the Bisbee is set at the base of the lowest (basal) conglomerate. The top of the formation was not seen in core. A few scattered fossils - gryphea and large gastropods- were found in the Bisbee, but none had any diagnostic value.

<u>Alteration Summary</u>. With the exception of that cored in DDH T3, the Bisbee is only weakly altered. Carbonate units show weak, patchy development of skarn, while the clastic units are weakly hornfelsed. Very weak silicification affects nearly all Bisbee lithologies.

In T5, short intervals of arkose and siltstone are sufficiently silicified to resemble "novaculite." Bisbee lithologies in T3 are slightly more altered than those in other holes, but overall the formation has the appearance of being only lightly "cooked", in contrast to the distinctly strong alteration found in the Paleozoic section.

Mineralization Summary. Aside from intervals of up to 10% disseminated pyrite in arkose, siltstone, and limestone, and the weak development of manganese oxide replacements in a few carbonate intervals, no significant mineralization was observed in the Bisbee. In the case of the limestone units, this is slightly surprising as these were the hosts for most of the known ore in the district. The only hint of the

possible proximity of ore in the Bisbee was found in T5. Here, euhedral blue green and green fluorite crystals line vugs in a brecciated, silicified limestone. The fluorite's character and nature of occurrence are identical to that which occurs adjacent to the large oxidized orebodies of the Benton-Albert stopes in the Empire Mine.

Structural Characteristics Summary. Most of the clastic units in the Bisbee show strong, brittle deformation in many sections. Numerous fault breccias, shears, and punky gouge zones are common in the clastic units, as are long sections of high fracture density, many of which are dominated by nearly vertical breaks. Carbonate units, contrastingly, generally have low fracture density and many have the appearance of having been smeared along very closely spaced shear planes. While no evidence of true plastic deformation was observed in the carbonates, the "smearing" phenomenon closely approximates it.

## Igneous Rocks Scheffelin Granodiorite

Distribution and Thickness. Two sills (?) of Scheffelin DDH T1 which also bottomed in granodiorite occur in The upper sill occurs in Epitaph limestones, granodiorite. The lower sill cuts Earp shales and and is seven feet thick. is three feet thick. The drill hole cut 11 feet of Scheffelin in the bottom before caving prevented further progress.

Lithology. The Scheffelin is a reddish gray, medium grained, equigranular rock composed of plagioclase, minor K feldspar, minor quartz, and accessary biotite and hornblende. Minor oxidized pyrite was observed in the Earp sills.

<u>Alteration Summary</u>. The granodiorite is moderately altered. Feldspars have been thoroughly kaolinized, while the mafic minerals have been, in part, chloritized. The rock is stained with iron oxides, either from oxidized pyrite or from the alteration of mafic accessory minerals. The limestones and shales immediately adjacent to the sills are not visibly altered - no contact metamorphic effects could be found despite a careful search. This is in accord with findings at similar exposures underground in the Lucky Cuss Mine, and suggests that the Scheffelin was intruded in a very dry condition at fairly low temperatures.

<u>Mineralization Summary</u>. The Scheffelin contains trace amounts of accessory pyrite and is cut by tiny pyrite-bearing veinlets. Given the unaltered and unmineralized nature of the immediately adjacent wall rock and the generally unmineralized nature of the Scheffelin, it is doubtful that the intrusive was responsible for the CRD mineralization in the Tombstone district.

## MINERALIZATION

All holes in the 1988 drilling program contained mineralization of one type or another, as did all of the formations cored in the holes. Mineralization varied greatly, however, as to type, mineralogy, and intensity, ranging from skarns containing disseminated pyrite, to thin, but intensely mineralized base and precious metal "mini-mantos", to massive, argentiferous manganese oxide replacements. This section classifies and describes the various types of occurrences found in holes T1-T5.

## Base Metal Replacements

This classification embraces bedding-controlled base metal replacements of unskarned limestones. Numerous occurrences are found in the core, but the best examples are found in T1 and T2 in the Epitaph and Colina, respectively.

The megascopic mineralogy of this type of Mineralogy. mineralization is rather simple, consisting of galena, sphalerite, chalcopyrite, and pyrite, along with possible tetrahedrite and argentite. Sphalerite is typically of the honey-yellow amberjack variety; blackjack is rare. Galena, chalcopyrite usually occur as coarse sphalerite, and interlocking grains, up to .5", while pyrite is more fine the other and disseminated among sulfides. grained Tetrahedrite (?) and argentite (?) are very fine grained and occur in tiny veinlets cutting other sulfides. Typically the

paragenesis is amberjack sphalerite - blackjack (if present) galena - chalcopyrite - tetrahedrite (?), with pyrite occurring throughout the sequence.

Form. This type of mineralization has extremely sharp footwall and hanging wall contacts which are coincident, or nearly so, with the bedding of the host limestone. In gross aspect the contacts are planar, but in detail are slightly irregular. Ranging from 0.1" to 2' in thickness, they resemble miniature mantos in most aspects.

<u>Host Rocks</u>. Base metal "mini-mantos" occur exclusively in "clean" recrystallized, but unskarned, limestone; most commonly that of the Epitaph or Colina. These replacements clearly avoid dolomite units, some having grown from a limestone bed up to the contact of a dolomite unit and stopping abruptly. Such replacements have not been observed in "dirty" limestones, or in any of the clastic units.

<u>Associated Alteration</u>. Most "mini-mantos" are unaccompanied by any discernible alteration, having razor sharp contacts against unaltered limestone. In a very few cases, manganese oxides, in the form of dendrites and/or pinkish limestone, form what may be a balo in both the hanging and footwall of the small mantos. This possible "halo mineralization" occurs several feet (up to eight feet in one case) above and below the manto and is separated from the manto by totally barren limestone.

Metal Values. Manto-type replacement mineralization found in the drill holes is usually of ore grade, but never approaches ore widths. Typical assays are as follows: 1 foot grading .02 opt Au, 5.81 opt Ag, .09% Cu, 8.3% Pb, and 7.5% Zn; 1 foot of a 2 foot thick interval grading 1.3 opt Ag, 1.04% Cu, 5.0% Pb, and 3.94% Zn; and a 1.5 inch thick interval grading 1.87 opt Ag, 2.14% Cu, 10.4% Pb, and 5.4% Zn. Most intervals of mineralization in this classification were too thin (.1"-.5") to have yielded a meaningful assay because of excessive dilution.

## Pyrite Replacements

Pyrite replacements occur ubiquitously throughout the core from all of the drill holes, and show great variation in form and intensity. Included in this classification are weak disseminations in megascopically unaltered limestones, to more intense disseminations in skarn, to semi-massive manto-type mineralization in limestone.

<u>Mineralogy</u>. In some of its occurrences, especially in skarns and manto-type mineralization, pyrite replacements are accompanied by base metal sulfides, but in the majority of cases pyrite occurs alone as attested by assay values discussed below.

Form. Pyrite replacements take a variety of forms. Most common are weak, fine grained disseminations of anhedral pyrite. These occur over intervals of less than an inch up to scores of feet in a wide variety of lithologies and with

several associated alteration types. Veinlet- and fracturecontrolled pyrite often accompanies the disseminations, and some disseminations are clearly rooted in more intensely mineralized fractures. Weak, fine grained pyrite replacements also occur associated with wavy stylolites, especially in moderately to strongly recrystallized limestones. At the other end of the replacement intensity scale are the mantotype semi-massive pyrite replacements. These take a form very similar to base metal "mini-mantos" discussed above. They are bedding parallel with sharp foot and hanging wall contacts, and range between a tenth of an inch up to two feet thick. They are composed of up to 60-70% anhedral pyrite in a guartzcalcite gangue.

Host Rocks. Pyrite is found in all lithologies cored. It is perhaps most common in fine grained clastic rocks; grained disseminations, shales, as fine and siltstones typically in the 1-2% range, but, in places, forming up to 10% In these host rocks some, and perhaps most of of the rock. the pyrite may be diagenetic. Similar pyrite mineralization also occurs in arkose, sandstone, and conglomerate where it is more clearly of replacement origin. Weak to moderate (up to 10%) pyrite disseminations also occur in limestones and In many cases it shows a preference for thin dolomites. carbonate units sandwiched between fine grained clastic beds, or for the upper parts of carbonate beds immediately beneath shale or siltstone horizons. Semi-massive manto-type pyrite

mineralization occurs exclusively in carbonates. Unlike the base metal manto-types, it occurs both in limestone and dolomite.

<u>Associated Alteration</u>. Pyrite replacement mineralization occurs most commonly in altered rock, but is also occurs in rocks that are not visibly altered - most commonly in limestone. Hornfelsed clastic sediments and skarned carbonates, almost without exception, contain varying amounts of pyrite. In these cases, it is usually in the form of weak, fine grained disseminations, although manto-style semi-massive pyrite occurs in weakly skarned Earp and Horguilla limestones and dolomite. In other cases, pyrite usually accompanies moderate to strong silicification, and zones of weak to strong argillization.

Metal Values. Only in rare instances was pyrite replacement mineralization developed intensely enough to warrant assay. Even where strongly developed, metal values in this type of mineralization are low. The best one foot interval of disseminated pyrite mineralization in shale (T2) ran ND Au, 1.03 ppm Ag, .01% Cu, .16% Pb, and .18% Zn, while the best one foot interval of manto-style pyrite mineralization in Earp limestone (T1) ran .41 ppm Au, 43.75 ppm Ag, ND Cu, .03% Pb, and .01% Zn.

## Skarn-Related Mineralization

Skarn-related mineralization is found in all of the Tombstone drill holes and in every carbonate-bearing formation cored during the program. In all instances the mineralization

was either too weak or too thin to even approach economic grades.

Mineralogy. As with other types of Tombstone mineralization, that related to skarns is, megascopically, quite In decreasing order of abundance, the sulfide simple. minerals are pyrite, blackjack sphalerite, galena, amberjack sphalerite, and chalcopyrite. Typically all sulfides are anhedral and fine grained; exceptionally sphalerite grains Amberjack obtain dimensions of one inch in diameter. sphalerite is relatively rare, and where present is invariably rimmed by blackjack. Galena too is rimmed and/or corroded by early phase blackjack. This suggests an of skarn mineralization consisting of low temperature iron-poor amber sphalerite and galena, followed by higher temperature, ironrich black sphalerite, pyrite, and minor chalcopyrite. Very rare green sphalerite occupies an unknown position in this varagenesis.

Two types of skarn-related base metal Form. mineralization are present. One consists of mineralization which forms an integral part of the skarn occurring as disseminations, veinlets, and bands, while the other appears to be manto-type mineralization with relatively narrow skarn alteration selvages. The first type of mineralization varies in intensity from sparse, fine grained disseminations in which pyrite, galena, and sphalerite for only 1-2% of the skarn. More intense mineralization is marked by the appearance of veinlets and narrow bands of sulfides making up 3-5% of the

intense skarn mineralization is marked by Most rock. irregular semi-massive "knots" of sulfides which locally make Such mineralization rarely exceeds up 10-15% of the skarn. intervals greater than 6". Chalcopyrite accompanies the other base metal sulfides only in this most intense style of The second, or manto-type skarn-related mineralization. mineralization is very similar to base metal "mini-mantos" described above, but have not been observed to exceed 1.5 inches in thickness. They differ, from those previously described, in having narrow selvages, one to three inches thick, of skarn minerals (garnet, vesuvianite, wollastonite, etc.); the entire assemblage occurring in unaltered or recrystallized limestone. Intergranular relationships suggest that, in this type of mineralization, a thin skarn manto first, with later base metal mineralization developed. replacing the central part of the skarn.

Host Rocks and Associated Alteration. Mineralized skarn in every carbonate bearing lithology in every found is formation cored during the program, but is best developed in the Earp. The Earp was probably most susceptible to skarning because of the "dirty", argillaceous nature of the limestones, and the numerous, thin clastic units interbedded throughout The Colina, being the "cleanest" the carbonate horizons. carbonate formation cored, has relatively few, thin skarn Where best developed in the Earp or other horizons. formations, skarn has completely replaced the rock leaving no residual carbonate. Such skarn is composed of guartz, tan or

green garnet, wollastonite, vesuvianite, and rare brown pyroxene, which occur in widely varying proportions. Retrograde (?) chlorite, epidote, actinolite, and clays are generally unimportant constituents. Most skarn zones encountered in the drilling, then, appear to be prograde, an observation in accord with the increasing temperature suggested by the paragenesis of skarn-related base metal mineralization.

Metal Values. Only the more massive skarn-related replacement zones were assayed. A one foot section of py>sl>gl in a strong Earp limestone skarn assayed 1.09 ppm Au, 13.42 opt Ag, .38% Cu, 5.1% Pb, and 3.66% Zn. Another similar zone in the Earp assayed .03 ppm Au, 1.15 opt Ag, .16% Cu, 1.79% Pb, and .09% Zn. A one foot section of Earp garnetvesuvianite skarn containing a two inch gl-cpy-py-sl vein assayed .14 ppm Au, 2.05 opt Ag, .33% Cu, 1.52% Pb, and .72% Zn. A 1.5 inch "mini-manto" with a skarn selvage in the Colina assayed over one foot (note the very large dilution) yielded the following: .01 ppm Au, 21.09 ppm Ag, .16% Cu, 2.52% Pb, and 1.73% Zn.

### Fracture-Controlled Mineralization

This classification embraces vein mineralization as well as fracture- fault-, and breccia-related mineralization. It does not include the skarn-related vein and veinlet-controlled sulfides discussed above.

While mineralization of this classification is very widespread in all of the Tombstone drill holes, it is volumentrically insignificant, especially in the case of base metal sulfides. It is important to point out, however, that mineralized fractures are high angle. Since the most Tombstone drill holes were vertical or nearly so, the abundance and importance of this type of mineralization may, accordingly, be underestimated.

Mineralogy. All of the base metal sulfides discussed above, along with pyrite and manganese oxide minerals occur as Significantly, fracture-controlled mineralization. though, fault and breccia zones, as well as intervals of unusually high fracture density, are almost exclusively mineralized by manganese oxide minerals and finely disseminated pyrite. Base sulfides do not occur in such zones. Galena, metal sphalerite, and chalcopyrite, as well as pyrite occur in veins and veinlets which are isolated from intervals of major structure, suggesting mineralized fractures may predate many of the major structures found in the core. Of possible significance, too, is the fact that base metal sulfides (as well as fluorite) most commonly occur in high angle ( $\geq 80^{\circ}$ ) fractures, and often leave coexisting (but not necessarily coeval) flatter fractures barren. Sphalerite is the most common base metal sulfide found in these fractures, followed by galena; chalcopyrite is uncommon. Where both amberjack and blackjack are present, the former is always rimmed by the latter, as is the case with skarn-related sphalerites. Only

the larger veins contain significant gangue minerals. Quartz and calcite are the most common; the former is most closely associated with base-and precious metal-bearing veins, while the latter is more commonly associated with pyrite- and/or manganese oxide-bearing veins. Rhodochrosite and fluorite are uncommon gangue minerals.

Mineralization in the classification occurs as Form. fracture fillings, ranging in thickness from paper thin up to .1 inch across, and in veins up to two inches thick. As noted above, most mineralized fractures and veins have high-angle orientations. Fracture fillings tend to be tightly filled with pyrite and/or base metal sulfides, while veins (>.1") are and contain clusters of euhedral gangue commonly vuggy minerals: guartz, calcite, or fluorite. Dustings of iron or manganese oxides typically coat gangue or sulfide minerals in such vugs. Crustified veins are apparently quite rare; only a few were found in core from T5. These are one to two inches across and consist of thick outer layer of brilliant pink, acicular rhodochrosite, followed by a band of dense white quartz, and are cored by irregular masses of amberjack and pyrite.

<u>Host Rocks</u>. Fracture-controlled mineralization occurs in all lithologies cored. Fracture-controlled pyrite and manganese oxides are very non-specific in there occurrence, showing no preference among the various carbonate and clastic lithologies present. Base metal sulfide-bearing fractures and veins, conversely, show a distinct preference for carbonate

units, but apparently none between limestone and dolomite. Only rarely are tiny sphalerite veinlets observed in siltstone or shale units.

<u>Associated Alteration</u>. Alteration haloes around veins and fracture fillings in the Tombstone core are remarkably rare, and when present are exceedingly thin - much thinner than the half-width of the associated vein. Fracture fillings usually have no associated alteration with the exception of those containing manganese oxides minerals. These rarely have thin selvages of manganese dendrites and weak argillic alteration. Larger veins occasionally have thin bleached haloes, and, in the case of quartz-rich veins, an irregular halo of weak silicification.

# Manganese Oxide Replacements

Two distinct types of manganese oxide mineralization are apparent in the Tombstone area. The first type embraces weak disseminations, fracture coatings, and dendrite development which is widespread in the upper parts of the drill holes and of more sparing occurrence deeper in the holes in highly fractured or faulted areas. This type is probably related to supergene processes. The second type, probably directly related to hydrothermal processes, are massive replacements of limestone by manganese oxides which occur in drill holes T1, T4, and T5. A third type, manganese oxides which form apparent haloes around base metal mantos have been discussed in a previous section.

The first, or "supergene" type is briefly described here considered further and will not be in the follow-on This type of manganese mineralization is found discussion. mainly in the upper parts of the drill holes: above 640 feet in T2, above 395 feet in T3, above 490 feet in T4, and above 545 feet in T5. These intervals correspond roughly to the modern water table. In addition, this type of mineralization occurs deeper in the drill holes, but only in areas of brecciation, faulting, or dense fracturing. These data imply that this type of mineralization is related to supergene processes or circulating groundwater. The nature of the mineralization, too, suggests supergene redistribution of manganese, as the oxides occur primarily as dendritic fracture coatings or as weak impregnations in porous lithologies such as sandstone or arkose. Massive manganese oxide replacements are considered below.

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Mineralogy. The major manganese oxide in massive replacements is a fine grained, black, sooty mineral which is probably a mixture of pyrolusite and psilomelane. A few of the replacements contain a compact, dense, heavy manganese mineral which may be manganite. X-ray diffraction studies would be needed to positively identify the minerals occurring in these deposits. Manganese-rich replacements occur in T4 which also contain wulfenite and mimetite. These are discussed in a separate section, but it should be borne in mind that they may be related to the purely manganese oxide replacements of the present discussion.

Manganese replacements are intervals of limestone Form. which have been replaced by between 20 and 100 percent manganese minerals. The foot and hanging wall are welldefined and usually sharp, although weak dendrite-type mineralization or pink limestone may extend several feet from Thicknesses of massive or semi-massive the massive interval. mineralization range from six inches to 6.5 feet. Where bedding in the host limestone can be determined, the foot and hanging wall of the manganese replacements appear to be This type of mineralization, then, parallel or nearly so. appears to take the form of relatively thin mantos.

<u>Host Rocks</u>. Manganese mantos have only been found in limestone hosts. In DDH T1, a 6.5 foot manto occurs in strongly recrystallized Colina limestone. In T4, five mantos, between one half and three foot thick, also occur in recrystallized Colina. In T5, a one foot thick semi-massive manto occurs in moderately recrystallized Horquilla limestone.

<u>Associated Alteration</u>. The manto in T1 shows no hanging wall alteration, but a one foot interval of weak to moderate sanding is present in the immediate footwall. Mantos in T4 are surrounded by up to 15 feet of weak manganese oxide dendrite development, and the T5 manto has a two to three foot thick halo of pink manganiferous limestone alteration surrounding it.

Metal Values. The mantos in T4 and T5 have generally low metal values. Assay ranges are as follows: .01-.07 ppm Au,

1.58-11.73 ppm Ag, 5 ppm - .01% Cu, 76 ppm - .33% Pb, and 98 ppm - .27% Zn.

## Base Metal Oxide Mineralization

Mineralization in this classification bears many similarities to the ore that was mined during the heydays of the Tombstone district. Similar in both mineralogical composition and in physical form, the base metal oxide intercepts from the present drilling program differ from the old bonanza ores in grade, thickness, and depth of occurrence. Only three intervals containing base metal oxides were encountered during the drilling; all are in DDH T-4.

The mineralogy of the "ore" and gangue Mineralogy. minerals in the three intercepts is remarkably similar, with a few exceptions as noted below. Of the "ore" minerals, the lead oxides are the most common. Mimetite  $(Pb_5(AsO_4)_3C1)$  is the most abundant, occurring as yellow-green, or lemon-yellow pseudo-hexagonal crystals up to .1" long, which line fractures or vugs. Wulfenite (PbMoO<sub>4</sub>) is almost as common. It occurs as thin, delicate brilliant orange to red-orange panes, frequently twinned, and reaching .5" in size. Cerrusite is probably present but is too fine grained to be identified. No discrete zinc oxide minerals (smithsonite, hemimorphite, etc) were identified in any of the three intervals, but assay data suggests their presence. Silver oxides, notably bromargyrite (AgBr), are apparently very rare. Sooty black manganese oxides are present in two of the three intercepts. One intercept contained a sky blue mineral tentatively identified

as aurichalcite  $[(Zn,Cu)_5(CO_3)_2(OH)_6]$ . Gangue minerals include quartz, calcite, and fluorite. Quartz forms tiny crystalline druses, or occurs as multiple generations of botryoidal crusts. Calcite is either crystalline (tiny scalenohedra) or massive, and ranges in color from black (manganiferous) to clear, to brilliant yellow (iron oxidebearing). Fluorite is abundant as a gangue mineral and occurs as coarse druses of perfectly euhedral white to clear cubes. Orange iron oxides and thin films of manganese oxides commonly coat all "ore" and gangue minerals in these assemblages.

Two intercepts, those at 1320' (5' thick) and Form. 1470' (2' thick) appear to have a manto-like form. Both foot and hanging wall contacts are sharp and well-defined. While bedding is difficult to discern in the strongly recrystallized host rock, the contacts of the mineralized intercepts are parallel to stylolites and, therefore, probably also to original bedding. Mineralization is clearly of a replacement nature; no vein-like characteristics are present. The margins of the intercepts are composed largely of manganese oxides in guartz-calcite matrix, juxtaposed against unaltered а limestone. The interiors are very vuggy; some open spaces approach one inch across and are lined with crystals or druses of ore and gangue minerals.

The third intercept, at 2845', has an entirely different form, appearing to be weakly mineralized cave fill, lying immediately beneath a 13 foot natural void. Because of its

unusual nature and its potential importance is future exploration, this occurrence is described, in detail, below:

- 2792-2798': recrystallized light gray, stylolitic limestone.
- 2798-2800': as above, but containing 1/2" vugs containing orange calcite crystals (dog tooth spar) and amorphous clots of manganese oxides. Several fractures at 20° to core axis, haloed by manganese oxide dendrites and finely disseminated pyrite are present.
- 2800-2822.5': recrystallized white limestone with numerous thin seams of talcy clay at 30° to core axis. Specks of manganese oxides and orange carbonate scattered throughout.
- 2822.5-2823': white, slightly gritty, weakly sanded, recrystallized limestone.
- 2823-2836': void; possible oxidation cave.
- 2836-2836.5': very vuggy limestone. Vugs lined with perfectly clear rhombs of iceland spar; accompanied by white, sectile clay.
- 2836.5-2837': pulverent orangish-gray clay with calcite crystals scattered throughout.
- 2837-2842': very porous and vuggy limestone; vugs lined with calcite and guartz druses and with gray clay; numerous guartz and calcite veinlets. Strong iron oxide Liesegang banding is offset by 80° microfaults.
- 2842-2844': as above, but contains angular 1" fragments of argillized siltstone. Interval suggestive of a collapse breccia.

2844-2852': assayed interval. Very similar to unit immediately above. Vugs filled with pulverent bright orange iron oxides, and calcite crystals (scalenohedra with growth patterns outlined by iron oxides). Brilliant red-orange wulfenite panes and doubly terminated yellow-green mimetite crystals form druses in vugs and along fractures. Possible trace cerrusite and bromargyrite along open fractures.

2852-2865': recrystallized stylolitic limestone. Upper contact sharp; upper one foot contains only very rare tiny vugs and thin clay seams.

The base metal oxide intervals at 1320' and Host Rocks. in moderately recrystallized Colina limestone, 1470' occur moderately to strongly occurs in 2845' at while that recrystallized Earp limestone. In all three cases, the marble is clean, save for the presence of organic-rich stylolites. The lack of a silty or argillaceous component is indicated by the total absence of skarn minerals in either the hanging or footwall limestones of the mineralized intervals.

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The alteration in the 2845' Associated Alteration. 1320' The adequately described above. is occurrence occurrence shows a weak halo of manganese oxide, both as dendrites and fracture-fillings extending 13 feet into the feet into the footwall where it is and 13 hanging wall accompanied by minor pink manganiferous limestone. The 1470' occurrence shows a similar, but weaker manganese halo. It is accompanied, however by a zone of high angle fractures lined with fluorite, which extends nine feet into the hanging wall.

Metal Values. The best one foot of the 1320' occurrence ran .14 ppm Au, 32.33 opt Ag, .09% Cu, 1.27% Pb, and 1.60% Zn,

while the best one foot of the 1470' zone assayed .07 ppm Au, 2.84 opt Ag, .02% Cu, 2.8% Pb, and .21% Zn. The best one foot assay interval in the 2845' occurrence ran .07 ppm Au, .32 opt Ag, .01% Cu, .21% Pb, and .48% Zn. The reader is referred to the T5 log for complete assay information. It is noteworthy that assays of the complete intervals show that in these oxidized mineral occurrences the values for lead and zinc are almost equal and that zinc values, in places, exceed those for lead. Unoxidized base metal "mantos" at Tombstone usually show nearly equal lead and zinc values, or lead exceeding Insofar as zinc, in an oxidizing environment is far zinc. more mobile than lead, it is usually the case that oxidized mantos (at, for example, Leadville and Santa Eulalia) are strongly depleted in zinc by descending ground water. The values shown by the above three intercepts, accordingly, suggest in situ oxidation with little or no transport of the metal components away from their initial site of deposition.

X

#### STRUCTURAL GEOLOGY

the major thrusts of the Tombstone drilling One of program was to explore for manto mineralization in ground adjacent to major structures rather than to specifically target chimney-type mineralization within the structures themselves. The structures that were cut by certain drill holes were the result, then, of needing to explore footwall and hanging wall rocks, not of targeting the structure per se. The result of this approach was that little new information was generated concerning specific structures (i.e., the Lucky Cuss fault zone, etc). In some cases (e.g. T5) the nature and geometry of certain target-related structures became even less clear as the result of drilling. On the other hand, the wide scatter of hole locations generated some unexpected inferred data concerning major structures heretofore unknown. The following sections describe both the structural data directly observed in the drill holes, and those structures which drilling data imply to exist.

## Observed Structural Data

Lucky Cuss Fault Zone. DDHs T1, T3, and T4 were designed to explore ground in the hanging and footwalls of the Lucky Cuss fault zone. All were vertical holes of depths judged sufficient to cross the fault where it cut suitable carbonate host rocks deep in the Paleozoic section.

Drill hole T1 was collared 900 feet due East of the Lucky Cuss headframe where the fault zone strikes nearly north-

south. Data gathered from the accessible mine workings, as well as from published cross sections, indicates that the Lucky Cuss fault (between the surface and 400 feet underground) dips eastward at 70°. The planned fault intercept was 2,460 feet. It is most likely that the Lucky Cuss fault was actually cut at 2,655 feet, indicating on overall surface to intercept dip of 72°. At this depth, the fault is represented by a three foot calcite-cemented breccia zone in a nine foot thick shale sandwiched between two massive Earo limestone The breccia contains only weak, beds. disseminated pyrite mineralization. Weak to moderate galenasphalerite replacements occur, however, in the hanging wall limestone immediately above the fault.

Drill hole T3 was collared 450 feet east and 100 feet south of a large glory hole resulting from the mining of a 50 foot diameter chimney of argentiferous manganese oxide ore localized on the Luck Cuss fault. The fault in this area strikes almost due north and underground data indicated a southward rake of between  $85^{\circ}$  and near vertical. Measured surface and underground dips ranged from vertical to as low as  $55^{\circ}$  E. The planned fault/chimney intercept, assuming a  $80^{\circ}$ E dip and  $87^{\circ}$  rake, was 2,480 feet. The fault was actually cut at 2,426 feet, indicating an actual overall dip of  $79^{\circ}$ E. Several fractures, measured above and below the fault, dip between  $70^{\circ}$  and  $80^{\circ}$  which generally agree with the projected fault dip. The fault zone itself is seven feet thick and consists of sheared and brecciated, calcite- and guartz-

cemented arkose which contains several intervals of strong manganese oxide mineralization both as replacements and as vug and fracture fillings. No base metal replacements occur in immediate hanging or footwall limestones. The targeted chimney either dies out above the intercept depth, or was missed by the drill hole because of a change in rake.

Hole T4 targeted the down dip extension of mineralization at the Luck Sure Mine. The hole was collared 600 feet east of the mine openings where the Lucky Cuss fault still strikes almost due north before starting a sharp swing to the east (see Plate 1). Surface dips on various strands of the Lucky Cuss in the mine area range from 75°E to 85°E. Assuming an average dip of 80°E, the planned fault intercept was at 2,830 This point lies nearly in the center of the 13 foot feet. intersected between 2,823 and 2,836 feet. natural cave Fracture angles in rocks of the immediate foot and hanging wall tend to support an 80° fault dip. The nature of the mineralization associated with this intercept is described in detail in a previous section.

The intercept widths of the Lucky Cuss fault may appear surprisingly narrow. They are not uncharacteristically so, however, when compared to numerous segments of the Lucky Cuss observed both on the surface and underground where the width pinches and swells dramatically both along strike and down dip. Continuous exposures along several levels of the Lucky Cuss mine show the fault varies from a few inches thick in barren areas, up to perhaps ten or twenty feet in mineralized

areas (the exact width being obscured by replacement mineralization). Width may vary from one extreme to the other over a strike length of only a few feet. Accordingly, it is concluded that, although stratigraphic displacements could not be determined either underground or in core, the nature, dip, and intensity of faulting along the Lucky Cuss does not significantly change from the surface to the depths drilled.

Westside-Eastside Fissure. DDH T2 was collared near the intersection of the Westside-Eastside fissure, the Grand Central fault and the Tribute dike. Near the collar location, the dike dips west at 85° and the fissure dips 75° northwest (see Plate 1). Published cross sections and underground measurements in the Westside mine, however show the fissure to have an overall vertical to steep southeastward dip. It was anticipated, therefore, that the drill hole would pierce the dike-fissure intersection at depth, somewhere in the middle of Drilling results, however, showed the Naco section. otherwise. Neither the Tribute dike or the Eastside fissure were cut in the hole, indicating that both either maintain a vertical dip or reverse their indicated dip at depth. The presence of a thick section of Scherrer formation in T2 is undoubtedly due to faulting which may involve some movement on the Eastside fissure. It is probable, however, that pre-Bisbee faulting is also involved, as will be discussed in a subsequent section.

Empire Dike and Breccia Pipe (?). DDH T5 targeted mineralization associated with a strong breccia occurring at the intersection of the Empire dike and the Silver Thread Neither the dike or the roll crop out on the surface, roll. but both are mapped in detail on the 400 and 500 level of the Careful delineation of the breccia mass and Empire Mine. associated structures in the Benton-Albert stopes showed, on two levels, that the intersection of the vertical dike and roll axis had produced a vertically oriented breccia with a pipe-like geometry. T5, accordingly was targeted to intersect the breccia approximately 1,500 feet vertically below the breccia exposed on the 400 and 500 levels. As was the case with T2, the target was missed completely - neither the breccia or the dike were cut by the drill hole. The structural implications of the "miss" are as follows: 1) since the drill hole was oriented almost at right angles to the dike which was never cut, the dip of the dike must have changed from vertical to westward (away from the hole) at depth, and 2) the existence of the associated mineralized breccia at a depth of 2000 feet can neither be proved or disproved by this since the associated Empire dike was never drill hole penetrated, and there is no way of telling how closely the hole approached the dike along its course.

### Implied Structural Data

Perhaps the most important structural data generated by the program can not be observed directly in the core, but rather is implied, based on drill hole stratigraphic

information. In the following discussion reference is made to Plate 2 which shows the stratigraphic correlation between all of the Tombstone drill holes.

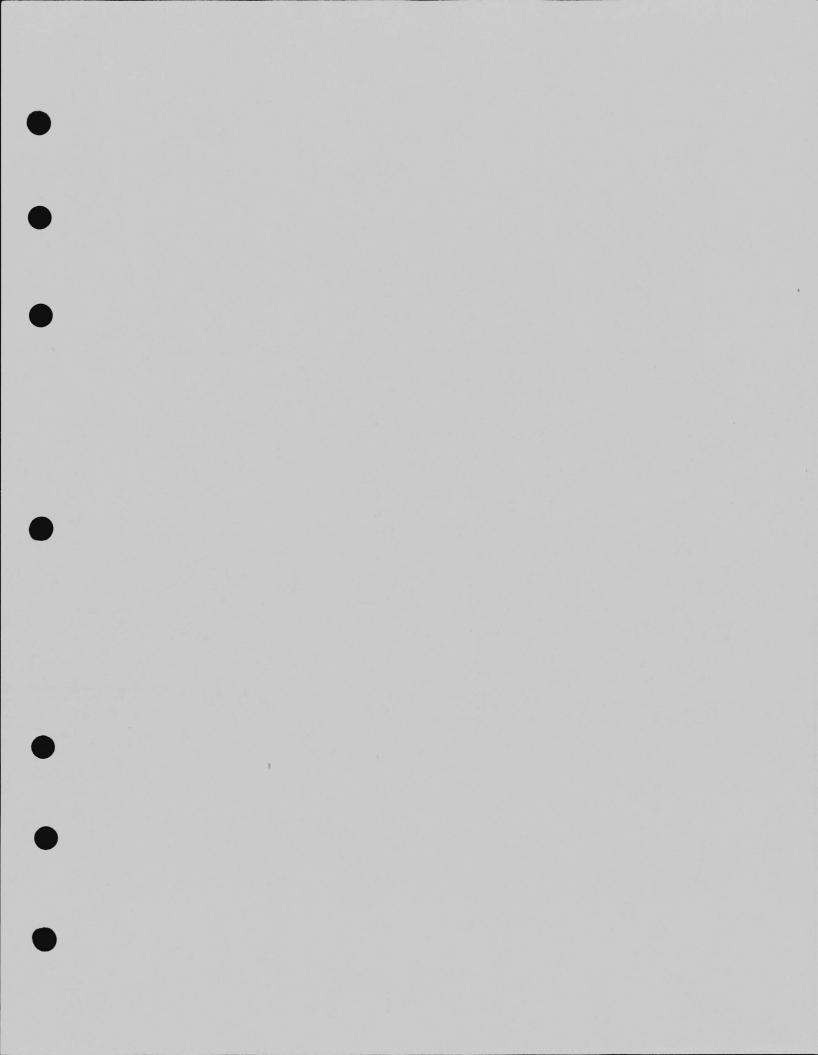
Inter-hole correlation clearly shows that the Paleozoic section has been affected by major faulting that is not expressed in the Bisbee and therefore must be pre-Bisbee in Evidence for this is found in the stratigraphy of holes age. In DDH T2 a thick section of Scherrer is present T2 and T5. immediately below the undisplaced Bisbee unconformity and a severely foreshortened section of Epitaph. above One solution to the geometry so imposed is shown on Plate 2 which suggests the Scherrer was faulted in on a generally eastdipping normal fault which also "decapitated" the upper parts The east dip is required to maintain the of the Epitaph. inter-hole stratigraphic integrity shown between T1, 3, and 4. In T2, the missing part of the Epitaph suggests a fault displacement exceeding 500 feet.

Diagonally across the Tombstone basin, DDH T5 revealed that the Horquilla limestone immediately underlies the Bisbee unconformity which, as in T2, shows no apparent displacement. The existence of the Horquilla in this location implies that the area containing holes T1-4 has been downdropped in excess of 2,300 feet relative to T5. Such a pre-Bisbee displacement may have occurred along one major fault as shown on Plate 2, or may represent the cumulative displacement of a series of west-dipping, west side-down step faults between T2 and T5.

Another noteworthy feature of Plate 2 is the apparent "flatness" of the Bisbee unconformity between drill holes T3, T1, T2, and T5. Closer spaced drilling might reveal local irregularities, but, in general, the unconformity would still maintain a fairly constant elevation. This implies that post-Bisbee faults (i.e. most of the faults within the Tombstone basin shown on Plate 1) did not produce significant vertical displacement. This is in accord with published mapping and the surface and on observations both personal with literature also suggests that lateral The underground. movement on such structures was minimal.

The only exception to this generalization, documented by the present drilling program, is the apparent movement on the Examination of Plate 2 shows that the Lucky Cuss fault. contacts between the various Paleozoic units in hole T3 are displaced sharply downward relative to those in T4, as is the Bisbee unconformity. Recognizing the distortions inherent in the diagram, and the presence of several cross faults (Plate 1), it still appears that these displacements are the result of drag, or successive, closely spaced stepwise displacements, caused by movement on the Lucky Cuss fault (compare Plates 1 and 2). If this is true, it can be further noted that the contact displacements decrease in magnitude upwards. Between T3 and T4, the Earp-Colina contact is displaced downward by 811 feet, the Colina-Epitaph contact by 640 feet, and the Epitaph-Bisbee unconformity by only 187 feet. These figures imply that the Lucky Cuss fault has been in existence at least

since Colina, and probably Earp time, and moved, perhaps as a growth fault, through Epitaph time. Movement continued or was reactivated in Bisbee time, offsetting the unconformity. If this analysis is correct, it marks the Lucky Cuss fault as one of the more important and fundamental elements of the structural fabric of the Tombstone district.



6200 Uptown Blvd. N.E., Suite 400 Box 27019 Albuquerque, New Mexico 87125 505/881-3050

May 1, 1989

#### FEDERAL EXPRESS

Mr. LaVerne Baxter Tombstone Development Company Route 2, Box 67 Grand Island, Nebraska 68803

> Tombstone Project Cochise County, Arizona File 3-033-000 1

Dear LaVerne:

Santa Fe will commence its 1989 drilling program on the Tombstone Development Company lease May 1st. This years drilling program will consist of one rig and will be as follows:

- T-6 will be drilled on the Luck Sure Patent at the same location as last years hole T-4. This hole will be drilled first.
- 2. T-7 will be drilled on the West Side Patent.
- 3. T-8 will be drilled on the Contention Patent.
- 4. T-9 will be drilled on the Silver Thread Patent at the same location as last years hole T-5.

I will supply you with a map of all drill sites prior to the completion of the drilling of T-6.

If you have any questions, please call.

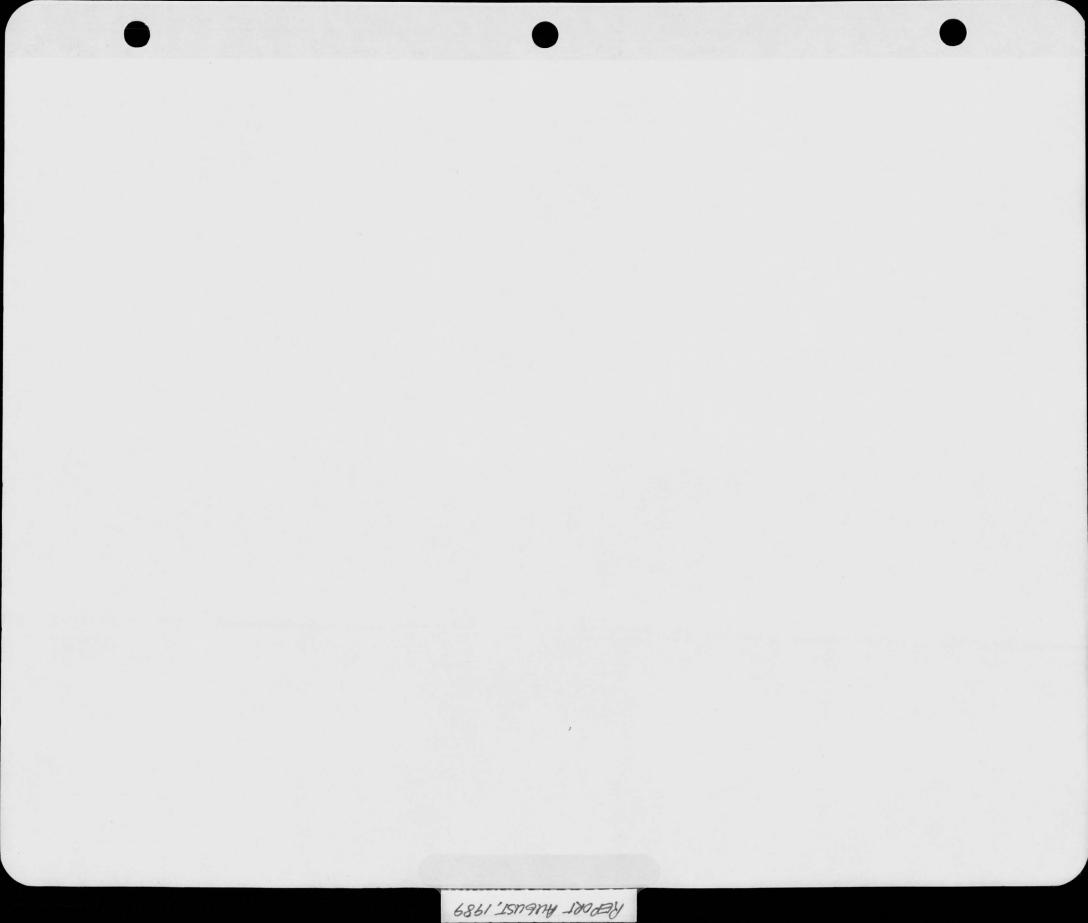
Very truly yours,

Dennis Cole Manager-Land

DC:dlu

Attachment

cc: W.D. Riesmeyer



Semi-Annual Report to Tombstone Development Company

on the

Tombstone Project, Cochise County, Arizona

W. D. Riesmeyer Santa Fe Pacific Mining, Inc. September, 1989

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6200 Uptown Blvd. N.E., Suite 400 Box 27019 Albuquerque, New Mexico 87125 505/881-3050

September 11, 1989

Re: Tombstone Project Cochise County, Arizona SFPM File 3-033-0001

Mr. LaVerne Baxter, President Tombstone Development Company P. O. Box 1445 Grand Island, Nebraska

Dear LaVerne:

Attached in accordance with the term of the Lease between Tombstone Development Company and Santa Fe Pacific Mining is the semi-annual report covering Santa Fe's work during this period of time.

If you have any questions concerning this report, please feel free to call either me or Duncan Riesmeyer of this office.

Very truly yours,

Dennis Cole Manager-Land

DC:fy

Attachment

cc: D. Riesmeyer

# Semi-Annual Report to Tombstone Development Company on the Tombstone Project, Cochise County, Arizona

#### Introduction

Santa Fe Pacific Mining, Inc. conducted diamond core drilling on the Tombstone Development Company lease between May 1 and August 7, 1989. Two deep core holes (T-6 and T-8), with an aggregate footage of 4,200', were completed during the period. Extremely difficult drilling conditions, with attendant increased costs, prevented completion of the anticipated four hole program.

The following report describes the results from the two drill holes. Included are descriptions of the stratigraphy, structure, and mineralization encountered in each hole. A geologic map showing the location of each hole, as well as graphic logs for T-6 and T-8 are appended as Plates 1 and 2, respectively. Complete drill logs and assay sheets form an appendix to the report.

#### Diamond Drill Hole T-6

DDH T-6 was drilled between May 1 and June 10, 1989, on an asimuth of S85W at -80 degrees from the collar of DDH T-4 (see prior report). T.D. was 1,900'. The bearing of the hole was a direct line between the T-4 collar and the main shaft of the Luck Sure workings, which exploited a chimney of argentiferous manganese and lead oxide ores, lying within the Lucky Cuss fault. Drill hole T-6 was sited to explore for similar chimney and attendant manto mineralization in and adjacent to the Lucky Cuss where the fault was projected to cut the Paleozoic limestones of the Earp formation, apparently 1700 feet below the surface. The proposed intercept laid on a straight line projection between the outcrop of the Luck Sure mineralization and the oxidation cave, and accompanying mineralization, cut in T-4 at a depth of 2840 feet, which is thought also to lie in the Lucky Cuss fault zone.

#### stratigraphy

DDH T-6 cut a section which included the Bisbee, Epitaph, Colina, and Earp formations. A complete stratigraphic description of these formations is included in a prior report; only a brief summary description will be given here.

Bisbee Formation. Rocks of the Cretaceous Bisbee formation were cored from the collar to 672 feet. The bulk of the unit consists of a generally coarse polylithic conglomerate, with clasts

in some sections reaching boulder size. Clasts consist primary of underlying Paleozoic units - limestone, marble, sandstone, siltstone, and shale, which are supported in a generally calcareous, sandy matrix. The second most abundant Bisbee lithology is fine to coarse sandstone, quartzite, and arkosic sandstone. The lower-most part of the Bisbee consists of siltstone and mudstone.

Epitaph Formation. In T-6, the Pennsylvanian-Permian Epitaph formation, consisting primarily of limestone and siltstone was cored between 672 and 1088 feet. The upper part of the unit is composed of fine siltstone and interbedded sandstones and arkose. The bulk of the unit is light grey limestone which has been recrystallized to a medium grained marble. The bottom of a lower siltstone unit marks the contact with the underlying Colina limestone.

Colina Limestone. The Colina limestone, cored between 1088 and 1521 feet, is also of Pennsylvanian-Permian age, and consists almost entirely of white to light grey, coarsely crystalline marble. Rare, thin dolomitic sections are present, as well as healed collapse breccia, and thin siltstone intervals. A small solution cavity and a four foot jasperoid interval were encountered in the upper part of the unit. The appearance of the first significant siltstone at 1521 feet marks the contact with the underlying Earp formation. A thin fault sliver of white marble between 1740 and 1750 feet is also tentatively assigned to the Colina.

Earp Formation. The Pennsylvanian-Permian Earp formation was cored between 1521 and 1738 feet and between 1750 and 1900 feet. It consists of limestone, calcite marble, and dolomite, along with interbedded siltstone, sandstone, and arkose. Coarse recrystallization of the limestones in the upper part of the unit make the distinction between the Colina and Earp carbonates difficult.

## Structure

Despite the number of faults and breaks mapped on the surface between the Luck Sure shaft and the collar of T-6, drilling showed very little structural disruption of the section except in the Bisbee, and in and near the Lucky Cuss fault. The contact between the Epitaph and Colina, and between the Colina and Earp formations can be readily correlated between holes T-4 and T-6, both dipping eastward between 45 deg. and 50 deg. This is in accord with bedding angles measured in both holes which generally range between 40 deg. and 60 deg. The numerous faults, noted on the accompanying drill logs, which cut the Paleozoic section, are generally high angle, but, as shown by hole to hole correlation, have only minor vertical offset.

A complication to this apparently simple structural picture is found in the attitudes of the Bisbee beds and the Bisbee-Epitaph contact. In both T-4 and T-6, bedding in the Bisbee dips very steeply - between 70 deg. and 90 deg., up to 50 deg. steeper than that of the underlying Paleozoic units. In addition, correlation between T-4 and T-6, show's the Bisbee - Epitaph contact to dip west at 6 deg. While this may be due to an extremely rugged unconformity surface, the attitude of the Bisbee beds suggest a tectonic cause. The shallow contact angle and the steep bedding angles suggest thrust faulting along the unconformity with possible overturning of the upper plate Bisbee. While there is some evidence of overturning reflected in the core, there is little or none of a thrust at the contact. For the time, the issue is unresolved.

The Lucky Cuss fault was intercepted in T-6 at 1738 feet, approximately 45 lateral feet from the projected intercept on the line between its outcrop at the Luck Sure and the T-4 intercept. Thus the Lucky Cuss dips 79 deg. between the surface and the T-6 intercept, and 75 deg. between T-6 and the oxidation cave in T-4. The cave and underlying mineralization encountered along the fault in T-6 suggest the Lucky Cuss zone to be about four feet thick, considerably thinner than the 20 foot thick section measured on the surface.

#### Mineralization

While several mineralized intercepts were obtained in T-6, none were of adequate grade and/or width to constitute ore. Of these, the three most significant are described below. Complete assay results for these and other assayed intervals are located in the appendix.

The interval 991.5'-992' contained a thin (0.5') replacement consisting of nearly massive blackjack sphalerite and minor disseminated galena. The intercept assayed 31.0 opt Ag, 0.81% Cu. 3.20% Pb, and 17.3% Zn. Mineralization appeared to be concordant with bedding and had razor-sharp contacts with the carbonate host. Mineralization occurred within an alteration shell of strongly recrystallized Epitaph limestone (marble). These characteristics suggest a typical CRD-type manto, albeit very thin. This "minimanto" may conceivably be rooted in the Lucky Cuss fault, 500 feet distant, up dip along bedding. It does not, however, extend further down dip to T-4 only 285 feet away.

The interval 1185.5'-1187' is a vuggy, strongly oxidized zone in Colina limestone in the immediate footwell of a barren three foot thick jasperoid. The zone contains manganese oxides lining vugs, as well as remnant kernels of galena (to 1" across), and scattered wulfenite. The zone assayed 3.28 opt Ag. 0.69% Pb, and 1.64% Zn. Projected 290 feet down dip along bedding to T-4, a silica-fluorite-pyrite-manganese zone is encountered which may be a manto-like correlative of the T-6 intercept.

The Lucky Cuss fault is probably represented by a two foot oxidation (?) cavern cut between 1738' - 1740', which is similar in nature to the cavern on the Lucky Cuss which was drilled in hole T-4. The T-6 cave (fault) has a siltstone hanging wall and a vuggy marble footwall. Unlike the T-4 cave, this cavern is floored only by vuggy, semi-massive manganese oxides; no wulfenite or anglesite was observed. Assays showed some lead oxides to be present, however, viz: 2 feet of 0.18 opt Ag, 0.06% Cu, 2.35% Pb, and 0.35% Zn. It is probable that the low Ag, Cu, and Zn values have resulted from migration of these metals away from this strongly oxidized zone, only the Pb remaining, fixed as oxidation zonestable, fine grained cerrusite.

Below the cavern, zones of manganese oxide impregnations and dendrites occur in Earp limestones, the strongest of which lies between 1797.5' and 1802'. Assays show the manganese oxides to be argentiferous, carrying 0.88 opt Ag; no other significant metal values are present. It is probable that this MnOx zone is an alteration/mineralization halo related to the Lucky Cuss zone.

## Diamond Drill Hole T-8

DDH T-8 was drilled between June 12 and August 7, 1989 from a site on the north end of the PBR Minerals open pit. The hole was drilled due west beneath the pit at an angle of -61 deg. T.D. was 2,300'. The hole was sited to explore the possible down dip extension of rich silver-gold ore associated with a generally north-trending complex of faults and the Empire-Contention dike. Published cross sections show bonanza-grade oxide ore was closely associated both with the margins of the dike and with the multitude of steeply dipping normal faults which progressively offset the dike westward. Because the best "shallow" ore was found where dikes or faults cut Bisbee limestones, it was believed that larger ore bodies might be found where the fault-dike system cut thick sections of Paleozoic carbonates at depth.

#### Stratigraphy

DDH T-8 cut a section consisting of the Bisbee formation and the Horquilla and Escabrosa limestones. These units are also thoroughly described in a prior semi-annual report, and only a brief summary of their characteristics is given below. Bisbee Formation. Unlike the Bisbee drilled in T-6 which contained coarse clastic sediments - sandstones an conglomerates, the 802.7 feet of Bisbee drilled in T-8 consists of fine clastics siltstone and mudstone, as well as limestone. Only rare, thin interbeds of sandstone or arkose were encountered in the section. The siltstones, in the upper 200 feet of the hole, are strongly shattered and altered, creating extremely bad drilling conditions. Drill progress, on occasion, was as little as one foot per 20 hour shift. Below the zone of shattering, the siltstones, black, tan, brown or red in color, are seen to have a variable sandy component and frequently grade into finer grained mudstones. Limestones showing varying degrees of argillization, dolomitization, and silicification, form interbeds, up to 50 feet thick, within the siltstone section.

The Pennsylvania-Permian Horquilla Horquilla Limestone. limestone forms the lower-most formation of the Naco group, underlying the Earp formation. The Horquilla was cored in T-6 between 802.7' and 1949.5'. The upper part of the unit is a thickly bedded, light grey to pinkish grey limestone which contains a few thin interbeds of siltstone. From 897' to 1332' the Horquilla is a limestone or dolomitic limestone breccia. Clasts average 2.5" to 3" across, and are supported in a calcite matrix. This sub-unit either represents a healed collapse breccia, or, more likely a tectonically shattered interval which was subsequently solution-healed with calcite. The lower half of the Horquilla in T-6 is composed of thickly bedded to massive, stylolitic, medium to coarse grained limestone which contains a few thick beds of dolomite and siltstone. Weak skarning affects the lower part of this limestone unit. The lower-most unit of the Horquilla is a thinly bedded dark grey to brown siltstone which disconformably overlies the Escabrosa.

Escabrosa Limestone. The Mississippian Escabrosa was cored between 1949.5' and T.D. at 2300'. The unit consists of thickly to massively bedded, moderately recrystallized, light to dark grey limestone, with lesser amounts of dolomite and dolomite limestone. Most parts have been affected by weak skarning and argillization. Stylolites are common throughout the section. Only one thin interbed of dark grey siltstone appears in the section.

Felsite Dike. A felsite dike intrudes the Escabrosa between 2179' and 2199'. The dike is an orangish cream aphaniticporphyritic rock with a sharp footwall contact, but a very diffuse hanging wall contact. Phenocrysts consist of small plagioclase and potassium feldspar crystals, rare rounded quartz "eyes", and blackish green biotite books. The aphanitic groundmass is very siliceous and has the appearance of potassic flooding. The rock probably has a quartz monzonite composition. The dike contains no sulfides, and shows no trace of argillization, unlike most other dikes seen on the surface or underground in the district. Structure. The lack of surface outcrops along the line of T-8, as well as poor pit exposures and the lack of underground and nearby drill hole data, make any structural analysis of T-8 extremely difficult and speculative. Despite these difficulties, the following observations are, nonetheless, offered.

Numerous faults were recognized on the T-8 core. Faults. Most have a dip ranging from 58 deg. to 78 deg.; regional structural data suggest that most, or all, dip eastward . Most show little or only moderate displacement, as differing lithologies or formations are not juxtaposed. The only apparent exception is a major fault very near the Bisbee-Horquilla contact. This fault, dipping eastward at 75 deg., when projected to the surface, corresponds closely to the location of the Contention fault as mapped by Butler. Because of its proximity to the formational contact, it is probable that the Bisbee has been down-dropped against the Horquilla along this structure. The northeasttrending Tranquility fault, approximately 600 feet north of the drill collar (see Plate 1), dips southeast and is probably responsible for down dropping the Bisbee-Horquilla contact approximately 245 feet relative to the position of the contact in DDH T-6 (see prior report).

As shown on Butler's cross sections, east-dipping faults progressively offset the contention dike to the west with increasing depth. No dike segments were encountered in T-8 with the exception of the above described felsite which occurs at a vertical depth of 1900'. Its projected position is 1000' west of the outcrop of the Contention dike as mapped by Butler. Drill data shows the felsite dike probably dips east at -60 deg. Its location and attitude militate against its being a segment of the Contention; it is more likely a segment of the Sulphuret dike which cutcrops just west of the vertical projection of the bottom of T-8. The absence of segments of the Contention dike further up-hole is, at present, unexplained.

Folds. Various tectonic fabrics observed in the core, as well as known structural trends elsewhere in the district, suggest that T-8 passed through the axis of one or more folds. Unfortunately, because fabrics in drill core cannot be absolutely oriented in three dimensional space, the true dip and strike of bedding, recorded on the drill log, cannot be determined. Accordingly, on the basis of one drill hole, positive proof of a fold structure cannot be obtained. However, it should be noted that the trace of the Girard "roll" (anticline) projects across the course of T-8 (Plate 1). Manipulation of measured bedding angles in the core suggest that the axis of the roll may have been crossed between 1000' and 1300'. This corresponds to the location of the intense breccia cored in the Horquilla, as described above. It may be speculated that this breccia formed as a result of shattering along the crest of the roll. That weak mineralization occurs associated with the bottom of this breccia unit may be analogous to roll crest replacements which are well documented elsewhere in the Tombstone district.

#### Mineralization

Mineralization occurs sporadically throughout the length of T-8. Most, however, is very weak. Only two intercepts yielded significant assay values, but those were not of sufficient grade and/or width to constitute ore. Assays for all intercepts chosen for analysis, in addition to the two discussed below, appear with the T-8 drill log in the appendix.

A three foot intercept, between 1648.6' and 1651.6', assayed 0.64 opt Ag, 0.62% Cu, 2.16% Pb, and 2.35% Zn. The mineralization consists of 20% coarsely disseminated sphalerite = chalcopyrite galena, hosted in strongly recrystallized and weakly skarned Horquilla limestone. Mineralization is accompanied by weak pink and black manganese oxides, as well as possible fine grained green garnet. In addition, the sulfides are surrounded by a halo of recrystallization up to 7 feet thick, and a shell of weak chloriteepidote skarn up to 23 feet wide. There is no obvious structural control for the mineralization, and it is probable that it is a part of a weak bedding controlled skarn-sulfide manto system of unknown source.

The best mineralized interval in DDH T-8 occurs associated with the felsite dike between 2154' and 2177.5', in the dike hanging wall, and between 2198.5 and 2199' in the footwall. The hanging wall zone consists of up to 30% total sulfides including galena, sphalerite, chalcopyrite and chalcocite. Trace amounts of various copper oxides and wulfenite are scattered through the section. The average tenor of the interval 2154' to 2177.5' (23.5') is .002 opt Au, 1.14 opt Ag, 0.61% Cu, 6.54% Pb, and 2.57% Reference is made to the appendix for a complete tabulation X Zn. of the metal values in this interval. The footwall intercept contains about 20% sulfides with a mineralogy similar to the section above. Associated alteration occurs mostly in the hanging wall and consists of weak recrystallization, and weak to moderate A one foot cavern, completely filled with massive sanding. crystalline calcite and dolomite occurs in the immediate footwall of the dike.

It will be noted that the felsite dike contains anomalous lead and zinc values only on its margins; the interior of the dike is completely barren. This, in addition to its megascopically unaltered nature and lack of even visible pyrite, suggests the dike

is post-mineral and was not the source of the surrounding CRD mineralization. Accordingly, it is suggested that mineralization is related to a pre-dike fault (the same as, for example, mineralization in the Westside fissure). Both the fault and mineralization are thought to have been subsequently intruded by the barren felsite.

#### SUMMARY

Two deep core holes were completed on the Tombstone Development Company lease as part of Santa Fe Pacific Mining's 1989 exploration program. A 1,900 foot hole was drilled in the Luck Sure mine area, and a 2,300 foot hole was cored under the PBR Minerals open pit. Both holes were targeted on major faults, the Lucky Cuss and the Contention fault complex, which hosted significant ore deposits in the near-surface environment. The holes were directed to intersect these structures at depth where they cut thick sections of Paleozoic carbonate rocks - locations holding potential for hosting large mantos and chimneys of CRD mineralization. Both holes cut the targeted structures, but found only minimal development of sulfide mineralization, the grade and/or thickness of which was inadequate to constitute ore.

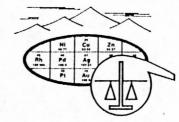
# APPENDIX

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Drill Logs and Assay Sheets for DDH T-6 and DDH T-8





SKYLINE LABS, INC. 1775 W. Sahuaro Dr. • P.O. Box 50106 Tucson, Arizona 85703 (602) 622-4836

REPORT OF ANALYSIS

JOB NO. UJV 086 June 13, 1989 T-6 991.5-992 PAGE 1 OF 1

SANTA FE PACIFIC MINING INC. Attn: Mr. Duncan Riesmeyer 6200 Uptown Blvd., #400 Albuquerque, NM 87125

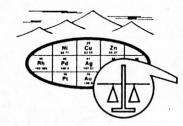
Analysis of 1 Drill Core Samples

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		FIRE	ASSAY				
ITEM	SAMPLE NUMBER	Au (oz/t)	Ag (oz/t)	Cu (%)	РЬ (%)	Zn (%)	
1	T-6 991.5-992	.002	31.00	.81	3.20	17.3	

cc: Mr. Fred Jenkins Mr. Dale Armstrong



James A. Martin Arizona Registered Assayer No. 11122



SKYLINE LABS, INC. 1775 W. Sahuaro Dr. • P.O. Box 50106 Tucson, Arizona 85703 (602) 622-4836

REPORT OF ANALYSIS

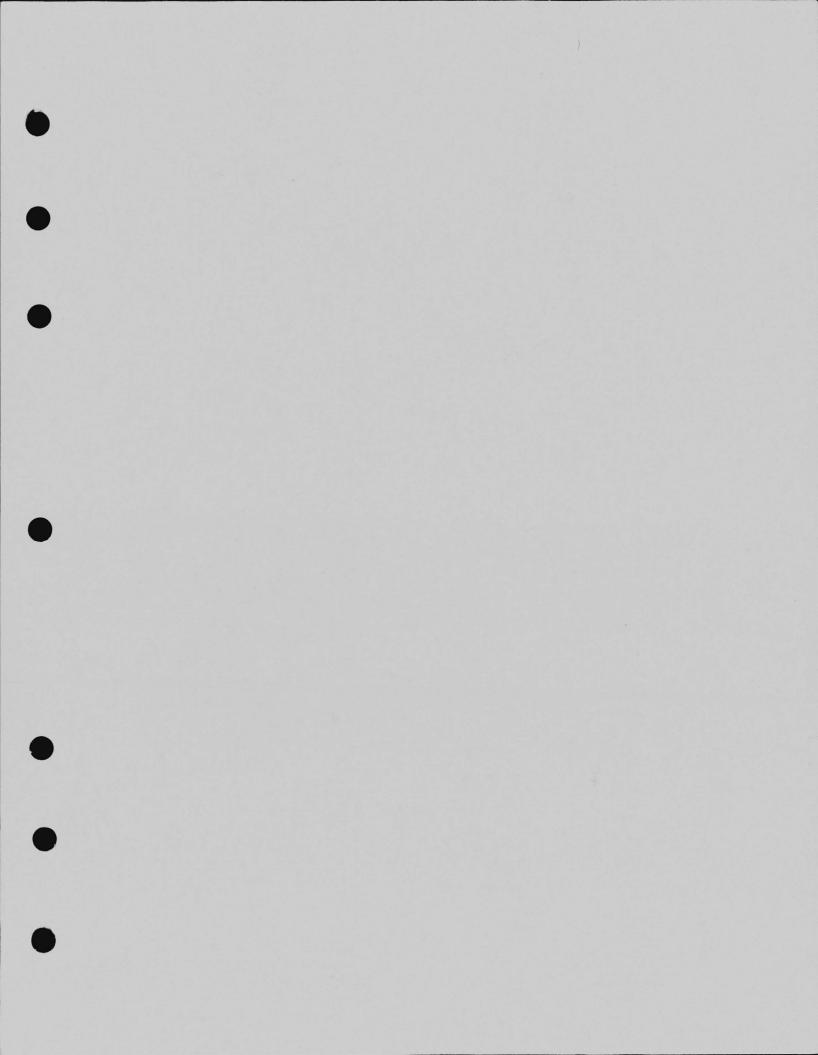
JOB NO. UJV 107 June 30, 1989 T-6 (1185.5 TO 1896) PAGE 1 OF 1

SANTA FE PACIFIC MINING INC. Attn: Mr. Duncan Riesmeyer 6200 Uptown Blvd., #400 Albuquerque, NM 87125

Analysis of 8 Drill Core Samples

		FIRE	ASSAY			
TEM	SAMPLE NUMBER	Au (oz/t)	Ag (oz/t)	Cu (%)	Pb (%)	Zn (%)
1	T-6 1185.5-1187.5	.010	3.28	.04	.69	1.64
2	T-6 1597.5-1602	<.002	.07	.01	.04	.09
.3	T-6 1741-1743	.008	.18	.06	2.35	.35
4	T-6 1770-1772.2	(.002	<.01	<.01	.02	.02
5	T-6 1779-1780.3	<.002	٢.01	٥.01	.01	(.02
6	T-6 1797.5-1800	(.002	.88	.02	.03	.32
7	T-6 1801.4-1802.4	4.002	.87	.01	.05	.18
8	T-6 1891.5-1896	(.002	.86	.02	.37	.35

cc: Mr. Fred Jenkins Mr. Dale Armstrong



DRILL HOLE		ALTERATION		DESCRIPTION	ASSAY DATA Page:
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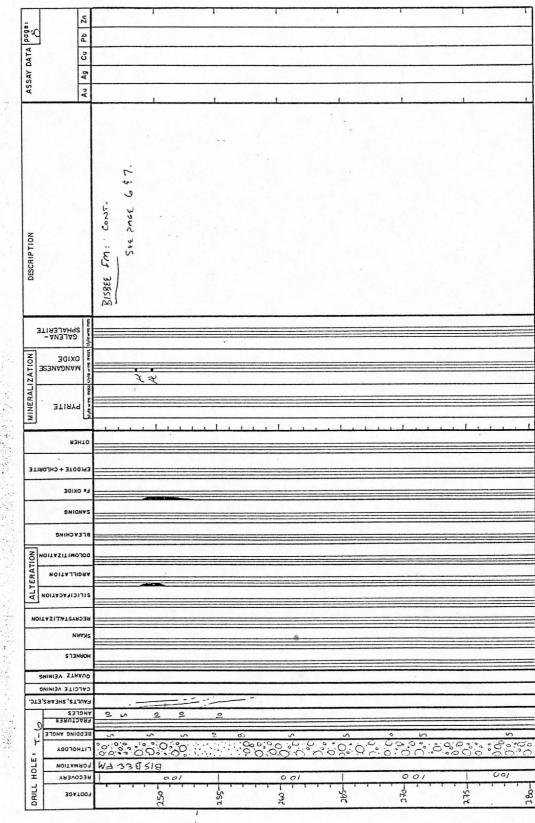
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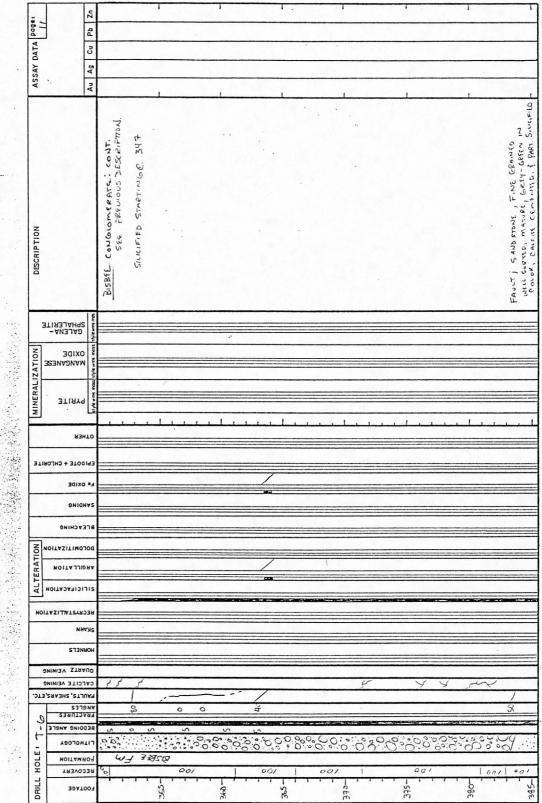
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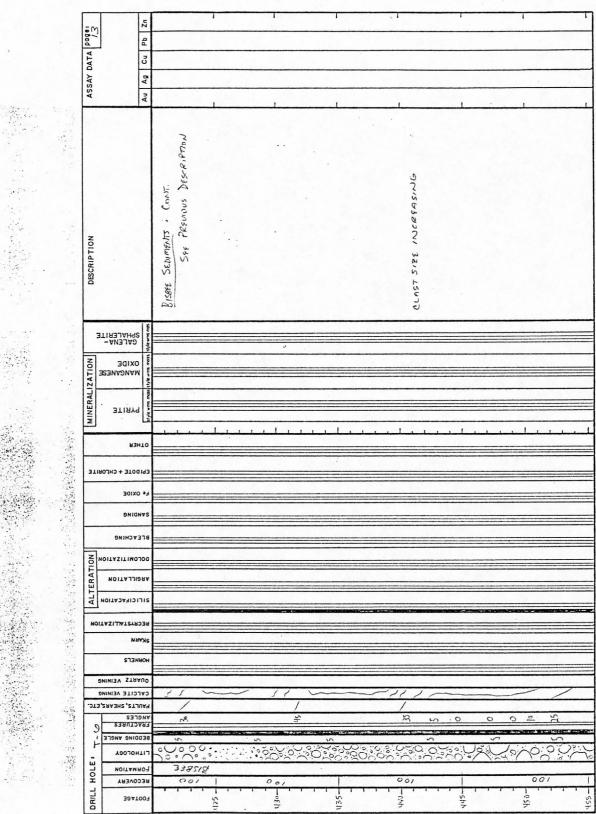
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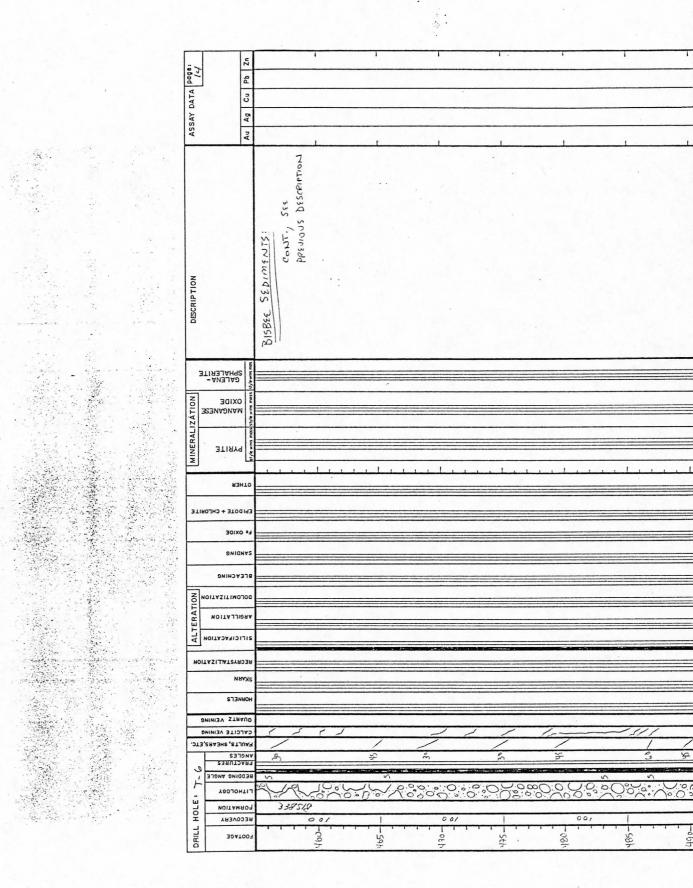
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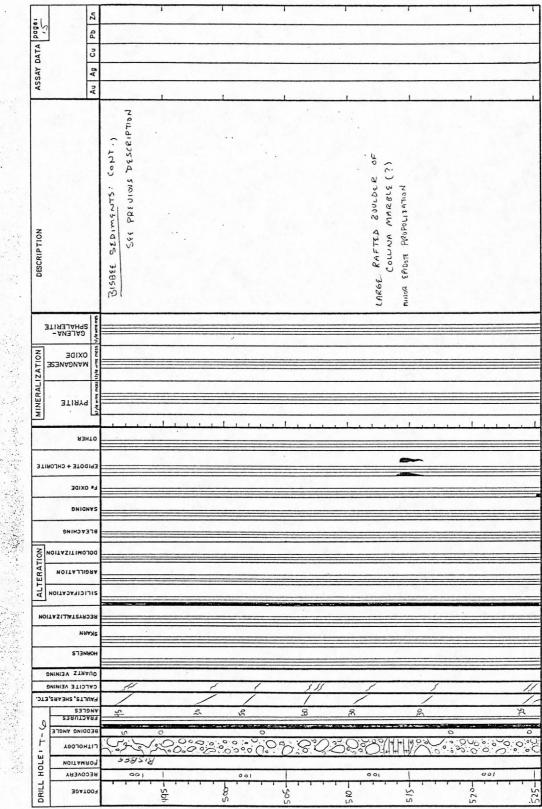
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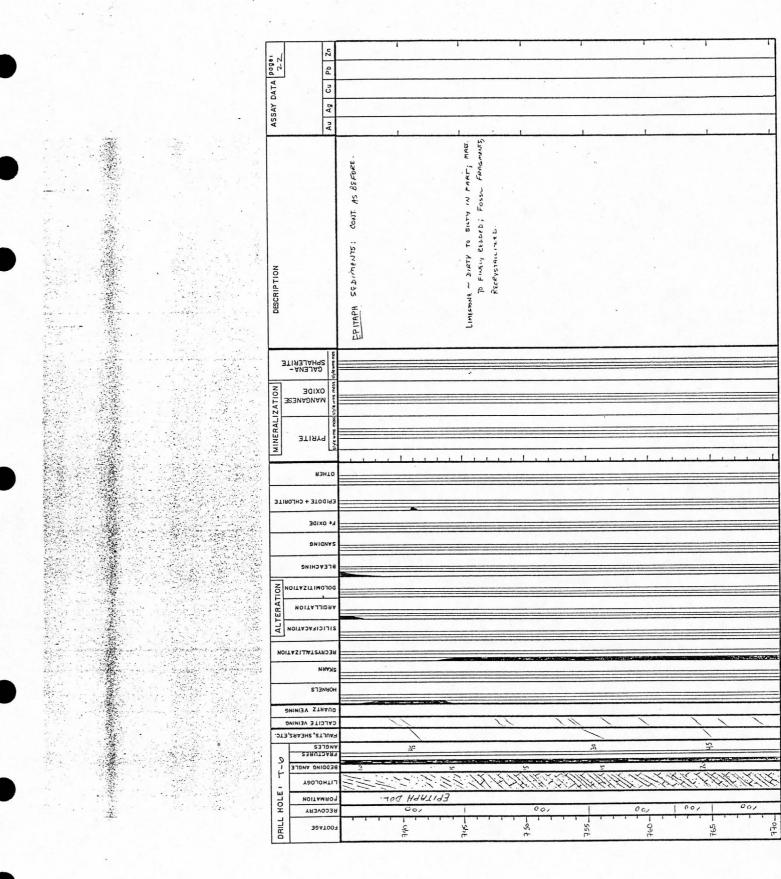
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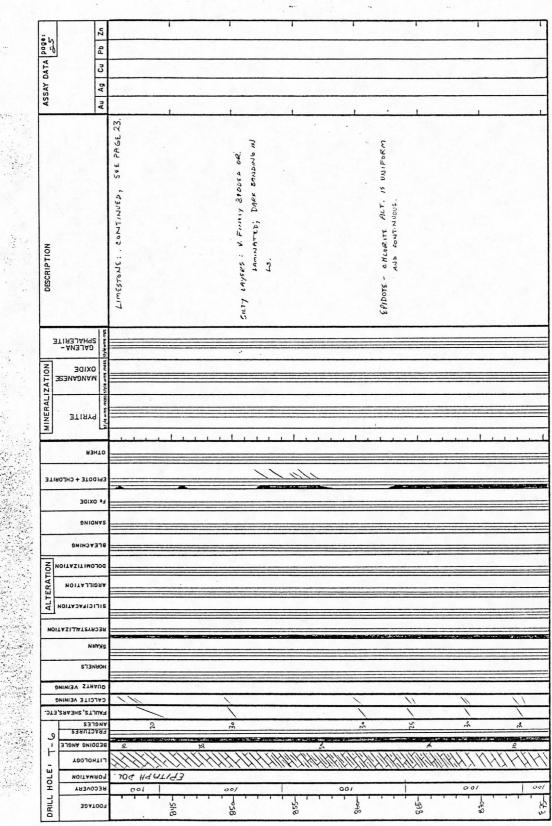
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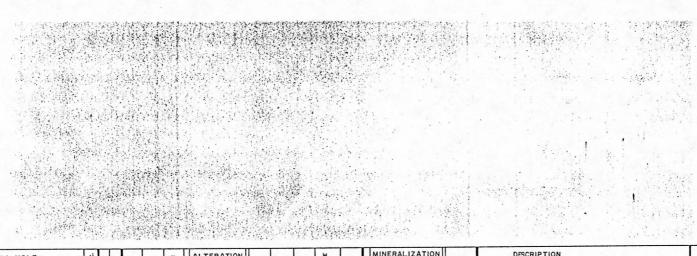
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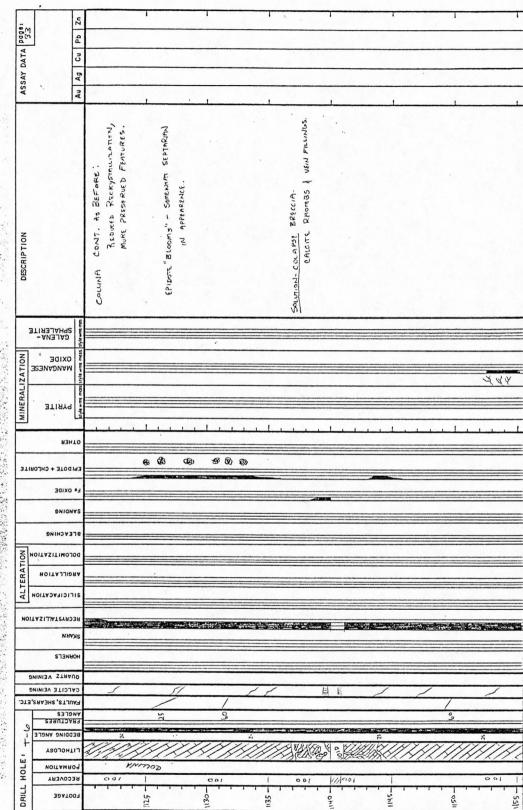
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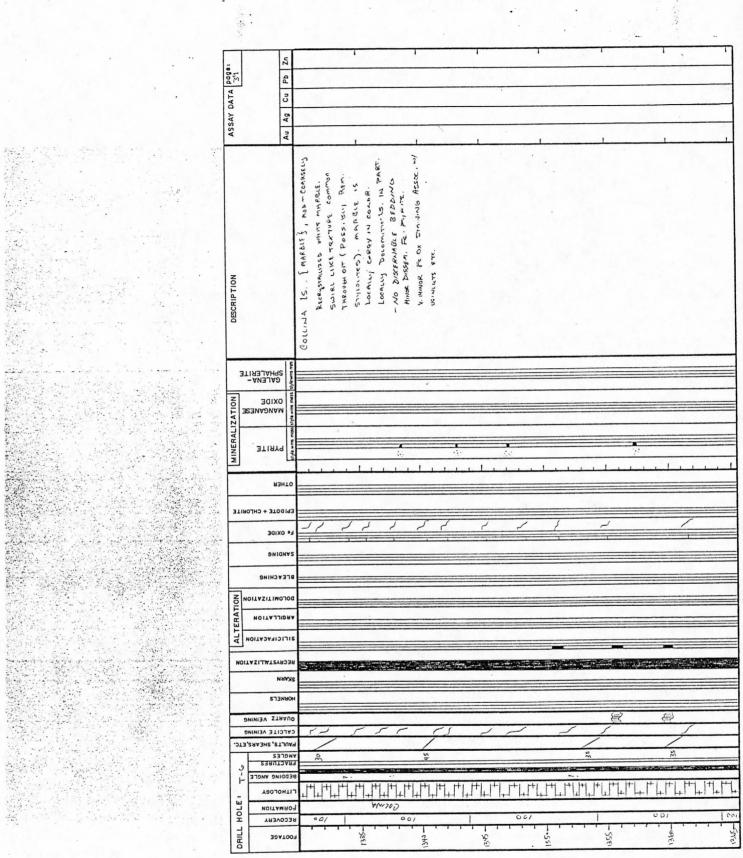
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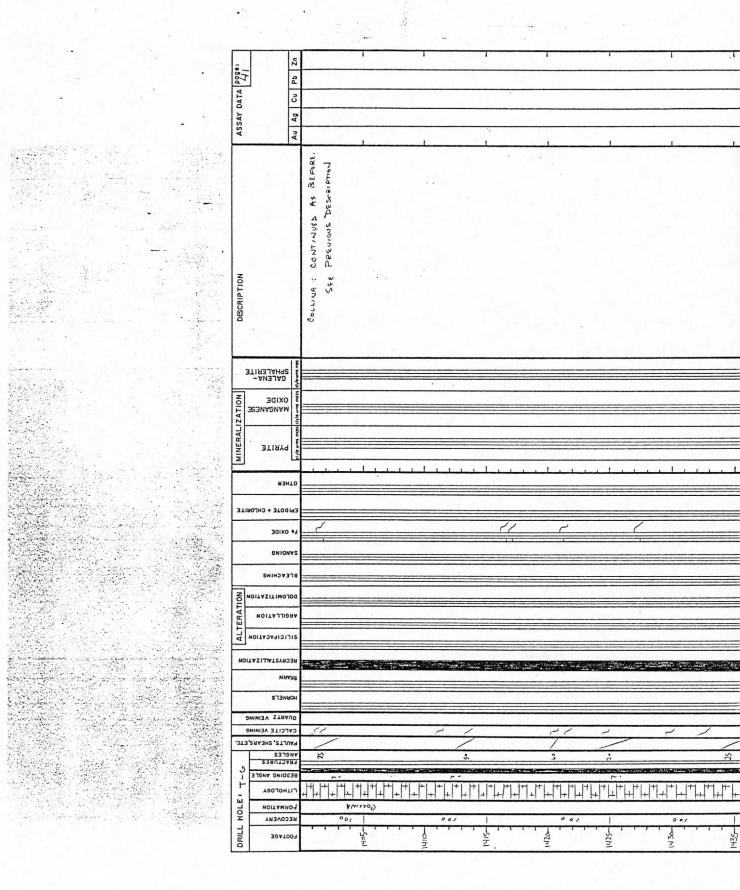
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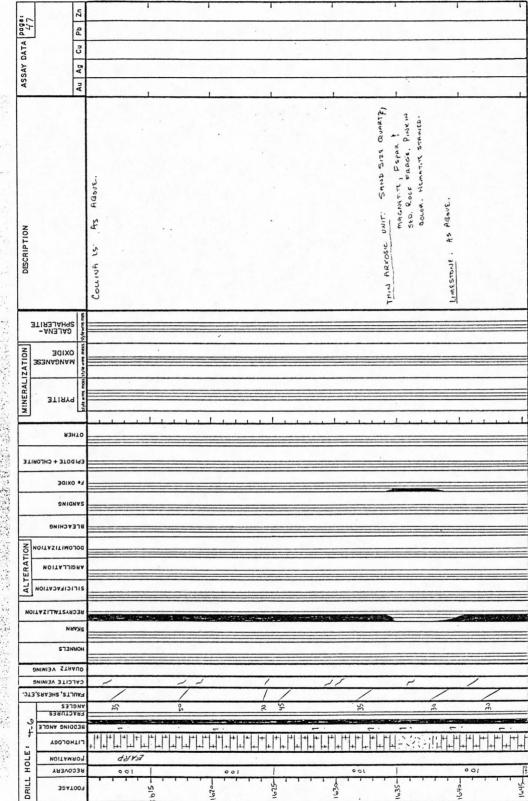
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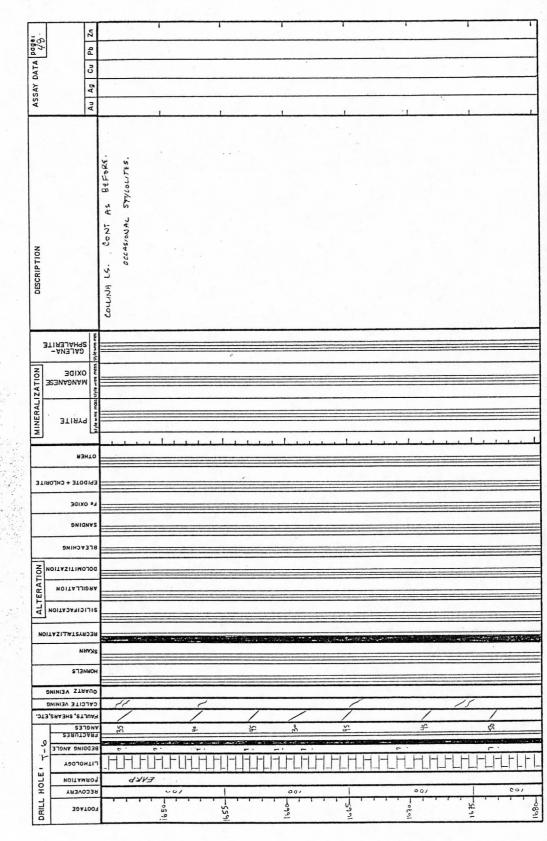
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FOOTAGE	FORMATION	BEDDING ANGLE	ANGLES FAULTS, SHE	CALCITE VEINING OUARTZ VEINING	HORNELS	SKARN	ECRYSTA	SILICIFACATION	DOLOMITIZATION	BLEACHING	SANDING	Fe OXIDE	PIDOTE 4	OTHER	PYRITE		MANG	SPHA	5		Au	Ag Cu	Pb
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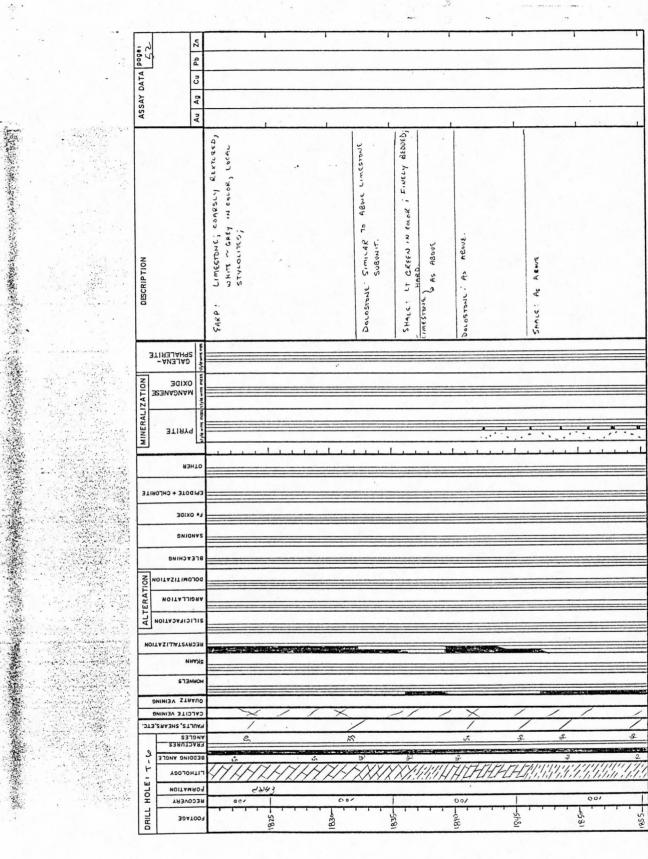
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			× × ·		Pyrine BLER. Rock is SICTY IN EART	SHATTERD ZONE. 1799'-		1787~1791	EARP FM: LIMESTONE; MASSIVE, STYLD RECRYSTALIZEL JOINTY, SLIPS A PLANE.	The stor	OXIDE	Έ
	-				Pyreite Scea. Rock is sictly in the	SHATTERED ZONE. 1799		1787~		mais style with mas	OXIDE	-

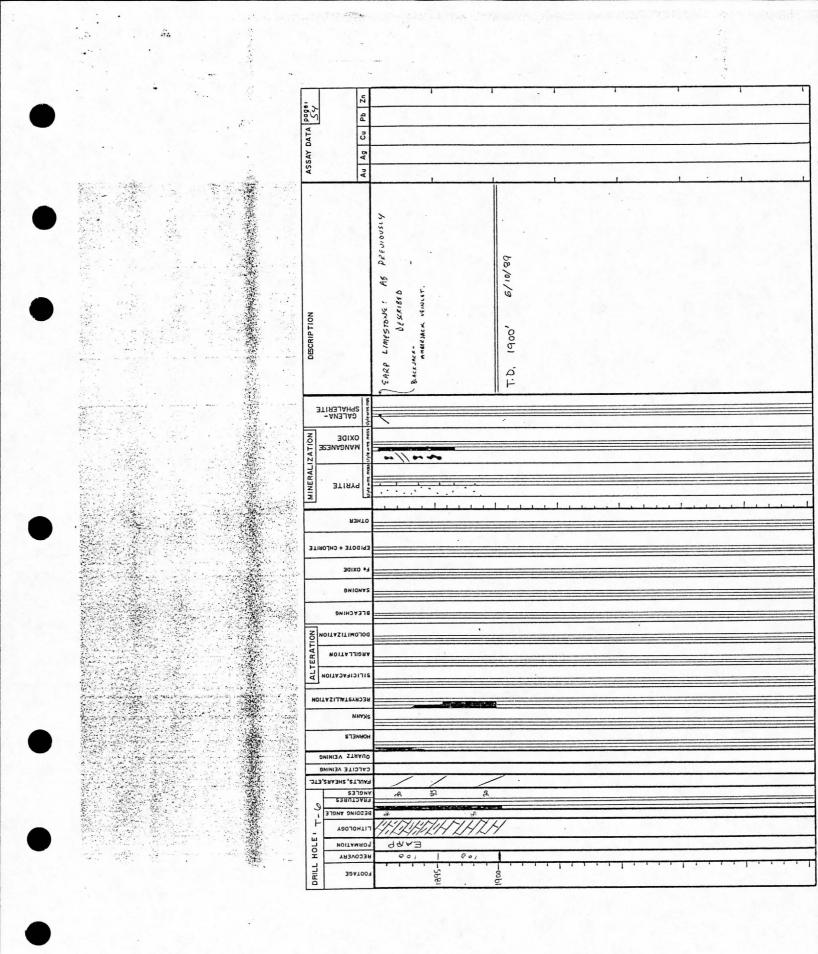
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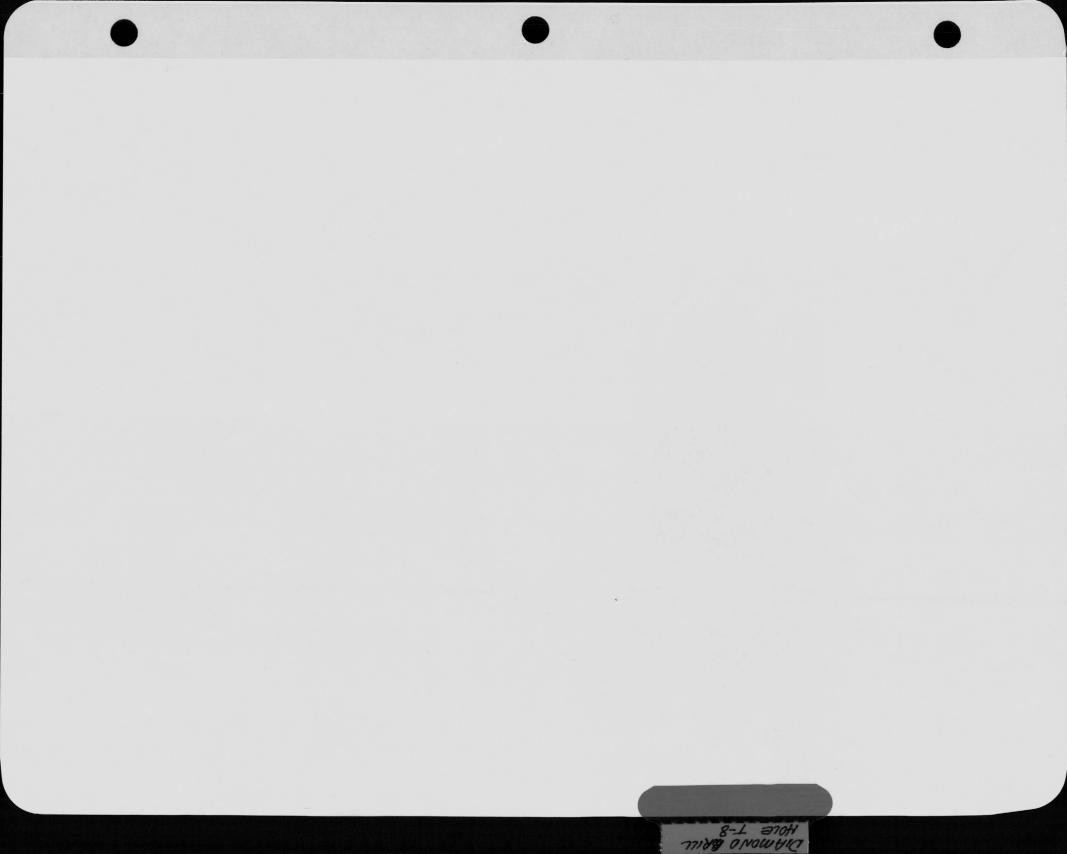


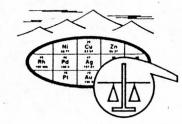
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FOOTAGE RECOVERY RECOVERY LITHOLOBY LITHOLOBY EEDDING ANGLE FOULTS, SNEJAR, SN	SKMM RECRYSTALLIZATION SILLICFACATION BILLIFACATION BULL	SANDING F4 OXIDE EPIDOTE + CHLORITE EPIDOTE + CHLORITE OTHER	PYRITE MANGANES MANGANES MANGANES SALENA- SALENA- SALENA- SALENA-	DESCRIPTION	ASSAY DATA Pagei 53
	34444       34444 <t< th=""><th>544 544 650 614</th><th></th><th>EARP: LIMESTONE: AS BEFORE DOLOSTONE: AS BEFORE SHALE: AS BEFORE LIMESTONE: AS BEFORE LIMESTONE: AS BEFORE EXCEPT SILTY IN PART</th><th>Au         Ag         Cu         Pb         Zn           -</th></t<>	544 544 650 614		EARP: LIMESTONE: AS BEFORE DOLOSTONE: AS BEFORE SHALE: AS BEFORE LIMESTONE: AS BEFORE LIMESTONE: AS BEFORE EXCEPT SILTY IN PART	Au         Ag         Cu         Pb         Zn           -







REPORT OF ANALYSIS

JOB NO. UJV 118 July 21, 1989 T-8 (230 TO 643.5) PAGE 1 OF 1

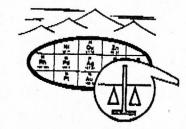
SANTA FE PACIFIC MINING INC. Attn: Mr. Duncan Riesmeyer 6200 Uptown Blvd., #400 Albuquerque, NM 87125

Analysis of 9 Drill Core Samples

		ETOE	ASSAY			
ITEM	SAMPLE NUMBER	Au	Ag (oz/t)	Cu (%)	Pb (%)	Zn (%)
_	T 0 000 005					
G	T-8 230-235 T-8 235-238	.006	(.01	<.01 <.01	<.01 <.01	.01
3	T-8 403-405	.004	.01	(.01	(.01	.01
4	T-8 405-410	(.002	(.01	(.01	(.01	(.01
5	T-8 525-529	.002	.10	.01	<.01	.04
6	T-8 529-533	.006	.22	<.01	<.01	.10
7	T-8 632.5-635	.006	.10	.01	<.01	.02
8	T-8 635-640	<.002	<.01	.01	(.01	.02
9	T-8 640-643.5	(.002	<.01	<.01	<.01	.01

cc: Mr. Dale Armstrong
cc: Mr. Fred Jenkins





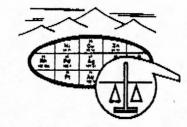
REPORT OF ANALYSIS

JOB NO. UJV 149 August 18, 1989 T-8 807-808.4 TO T-8 1669-1671 PAGE 1 OF 1

SANTA FE PACIFIC MINING INC. Attn: Mr. Duncan Riesmeyer 6200 Uptown Blvd., #400 Albuquerque, NM 87125

Analysis of 17 Drill Core Samples

	FIRE ASSAY											
ITEM	SAMPLE NUMBER	Au (oz/t)	Ag (oz/t)	Cu (%)	РЬ (%;)	Zn (%)						
1	T-8 706.3-707.3	(.002	.04									
ż	T-8 746.5-751.5	(.002	<.01	.01 <.01	(.01	.03						
3	T-8 807-808.4	(.002	(.01		<.01	.03						
4	T-8 813.5-817	(.002	(.01	(.01	(.01	.01						
5	T-8 1314-1319	(.002	<.01	(.01	<.01 <.01	.01						
6	T-8 1319-1324	(.002	(.01	( 01								
7	T-8 1324-1327	(.002	(.01	(.01	<.01	.01						
8	T-8 1390-1395	(.002	(.01	.01 (.01	.06	.06						
9	T-8 1476-1481	(.002	(.01	.01	.02 .02	.04						
10	T-8 1481-1486	(.002	<.01	<.01	.02	.01 (.01						
11	T-8 1486-1491	(.002	(.01	(.01	07							
12	T-8 1643-1644	(.002	(.01	(.01	.03	.01						
13	T-8 1646-1648.6	(.002	<.01	(.01	.02 (.01	.07						
14	T-8 1648.6-1651.6	.004	.64	.62		.02						
15	T-8 1651.6-1653.5	(.002	(.01	.01	2.16	2.35						
16	T-8 1653.5-1658.5	(.002	<.01	.05	07							
17	Т-В 1669-1671	(.002	.02	4.01	.02	.03 .03						



REPORT OF ANALYSIS

JOB ND. UJV 150 August 18, 1989 T-8 (2153 TO 2226) PAGE 1 OF 4

SANTA FE PACIFIC MINING INC. Attn: Mr. Duncan Riesmeyer 6200 Uptown Blvd., #400 Albuquerque, NM 87125

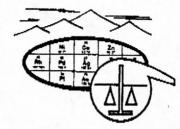
Analysis of 22 Drill Core Samples

ITEM	SAMPLE NUMBER	FIRE ASSAY Au Ag (oz/t) (oz/t)
1	T-8 2153-2154	<.002 <.01
the second s	T-8 2154-2155	(.002 3.15
	T-8 2155-2157	1.002 .59
4	T-8 2155-2158 **	.000** .00**
5	T-8 2157-2158	.002 1.96
6	T-8 2158-2161.5	.002 .68
7	T-8 2161.5-2162.5	4.002 2.14
8	T-8 2161.5-2163.5 **	
9	T-8 2162.5-2163.5	(.002 (.01
. 10	T-8 2163.5-2165	< <b>.</b> 002 <b>.</b> 30
11	T-8 2165-2167.5	.002 4.46
	T-8 2167.5-2172.5	<.002 .33 ·
	T-8 2172.5-2177.5	<.002 .56
14	T-8 2177.5-2182	(.002 (.01
. 15	T-8 2182-2186	(.002 (.01
16	T-8 2186-2187.5	<.002 <.01
	T-8 2189.5-2194.5	(.002 (.01
	T-8 2194.5-2198.5	(.002 .10
	T-8 2198.5-2199	(.002 10.29
20	T-8 2199-2202 *	.000* .00*
21	T-8 2202-2205	<.002 <.01
22	T-8 2205-2209	(.002 (.01
	T-8 2209-2211.5 *	.000* .00*
	T-8 2211.5-2216.5	(.002 (.01
25	T-8 2216.5-2219.5	(.002 (.01

Charles E. Thompson Arizona Registered Assayer No. 9427 William L. Lehmbeck Arizona Registered Assayer No. 9425 James A. Martin Arizona Registered Assayer No. 11122

		SKYLINE LABS, INC. 1775 W. Sahuaro Dr. • P.O. Box 5010 Tucson, Arizona 85703 (602) 622-4836	06		
				August	UJV 150 18, 1989 2 OF 4
	ITEM	SAMPLE NUMBER	Au	ASSAY Ag (oz/t)	
•	26 27	T-8 2219.5-2224 * T-8 2224-2226	.000 <del>x</del> (.002	.00 <del>x</del> <.01	<u> </u>

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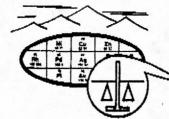
SKYLINE LABS, INC. 1775 W. Sahuaro Dr. • P.O. Box 50106

Tucson, Arizona 85703

(602) 622-4836

JOB NO. UJV 150 August 18, 1989 PAGE 3 OF 4 Cu Pb Zn ITEM SAMPLE NUMBER (%) (%) (%) 1 T-8 2153-2154 .02 .06 .10 2.28 2 T-8 2154-2155 7.35 2.45 T-8 2155-2157 3 .68 .00\*\* 2.62 3.05 T-8 2155-2158 \*\* 4 .00\*\* .00\*\* 5 T-8 2157-2158 1.33 5.90 4.25 6 T-8 2158-2161.5 .19 2.61 2.50 7 T-8 2161.5-2162.5 1.69 4.00 5.35 8 T-8 2161.5-2163.5 \*\* .00XX .00\*\* .00\*\* 9 T-8 2162.5-2163.5 .01 .08 .18 T-8 2163.5-2165 10 .12 1.60 .77 T-8 2165-2167.5 11 1.43 13.00 6.10 12 T-8 2167.5-2172.5 .56 .73 1.87 .08 13 T-8 2172.5-2177.5 1.31 1.50 14 T-8 2177.5-2182 .02 .07 .55 15 T-8 2182-2186 .01 .02 .02 16 T-8 2186-2189.5 .02 .05 .02 17 T-8 2189.5-2194.5 .02 .20 1.05 18 T-8 2194.5-2198.5 .02 .68 2.53 19 T-8 2198.5-2199 .49 15.60 2.60 20 T-8 2199-2202 \* .00× .00× .00¥ 21 T-8 2202-2205 .03 (.01 .04 22 T-8 2205-2209 (.01 .16 .04 23 T-8 2209-2211.5 \* .00× .00× .00× 24 T-8 2211.5-2216.5 (.01 .01 .01 25 T-8 2216.5-2219.5 <.01 .20 .07

Charles E. Thompson Arizona Registered Assayer No. 9427



-			Aug	NO. UJV 150 gust 18, 1989 PAGE 4 OF 4
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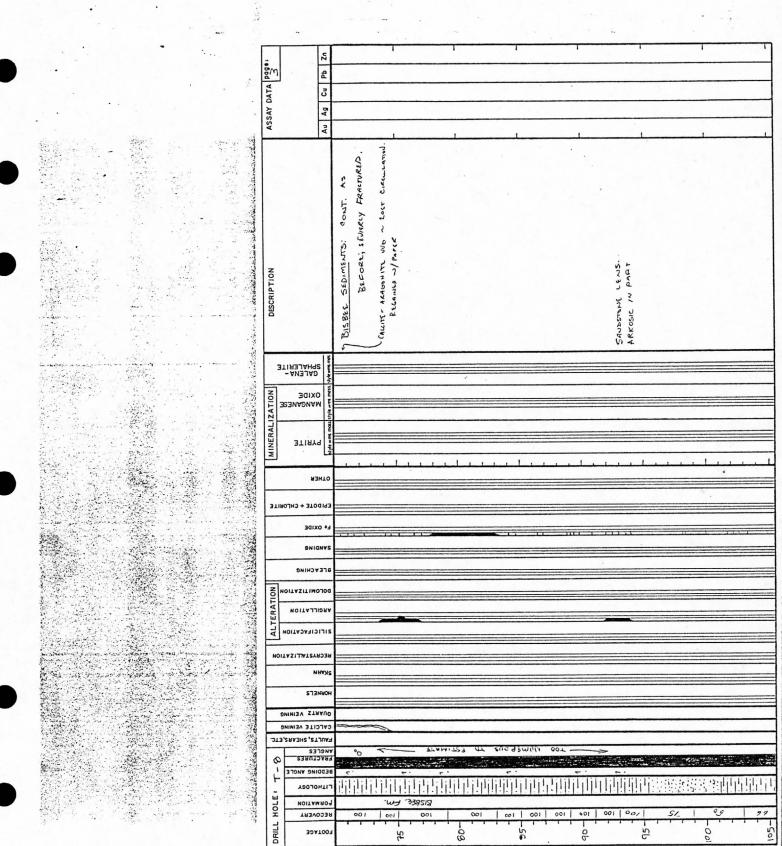
\*NOTE: No sample submitted.

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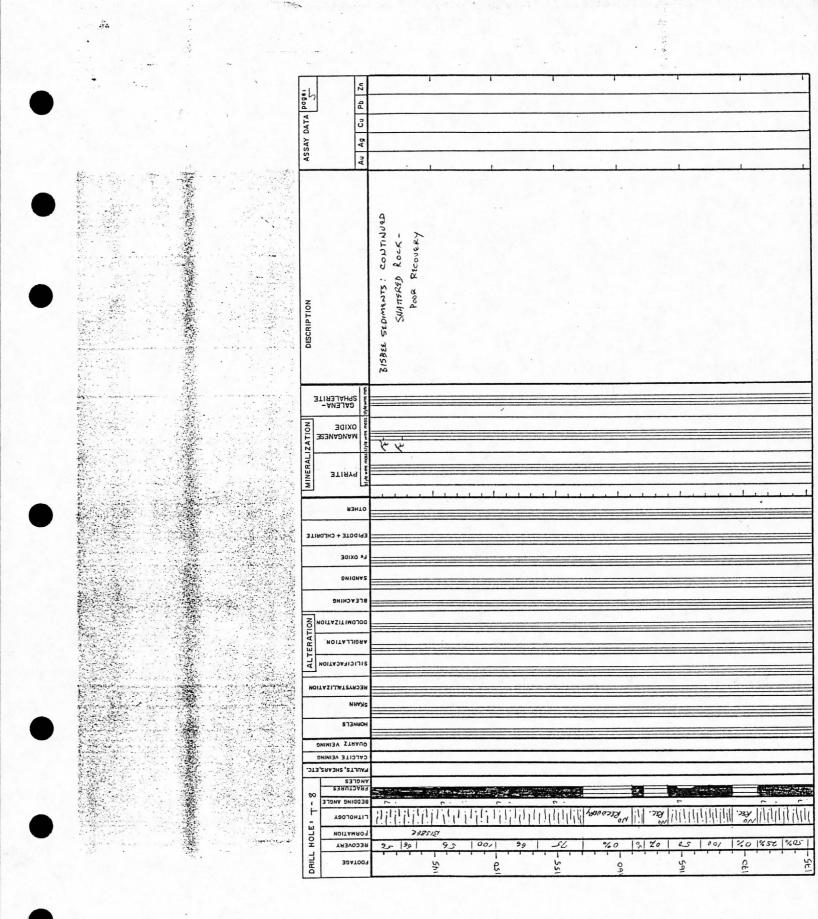


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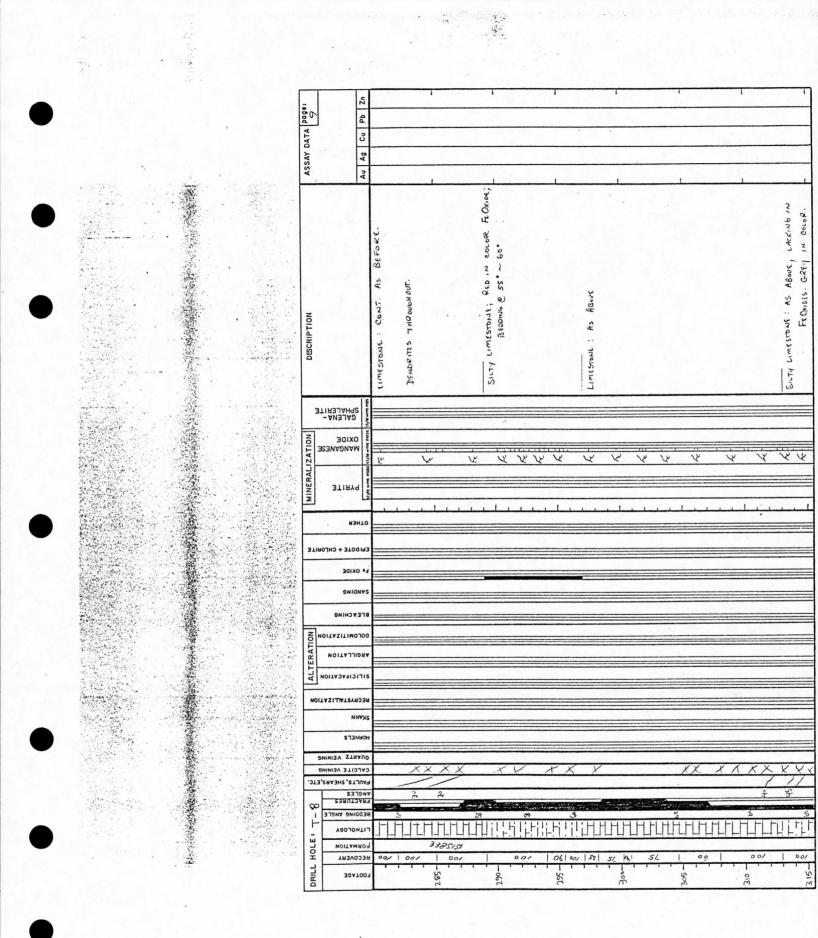
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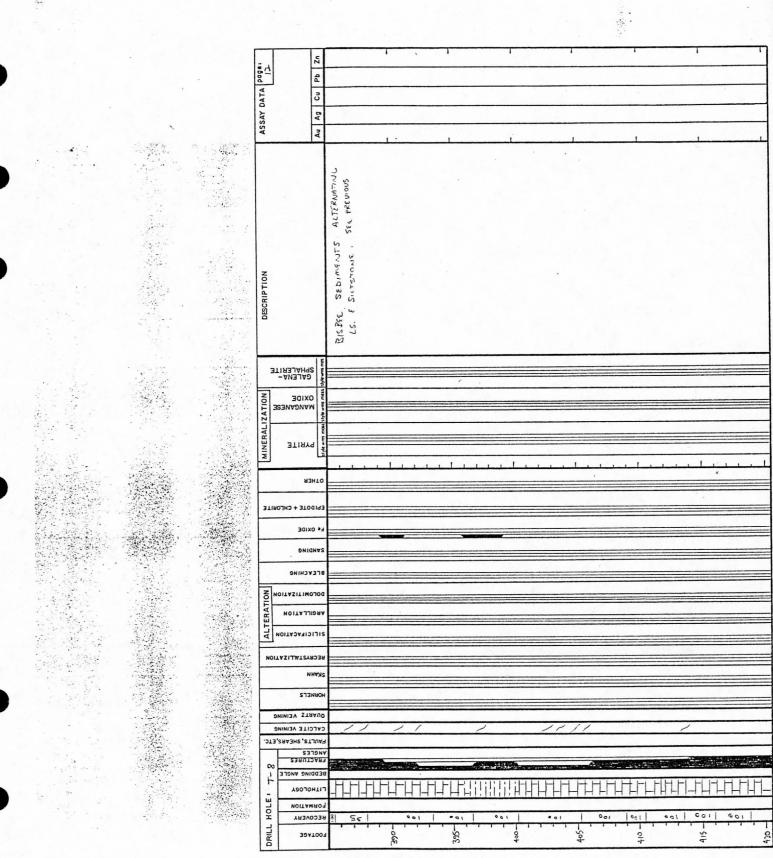
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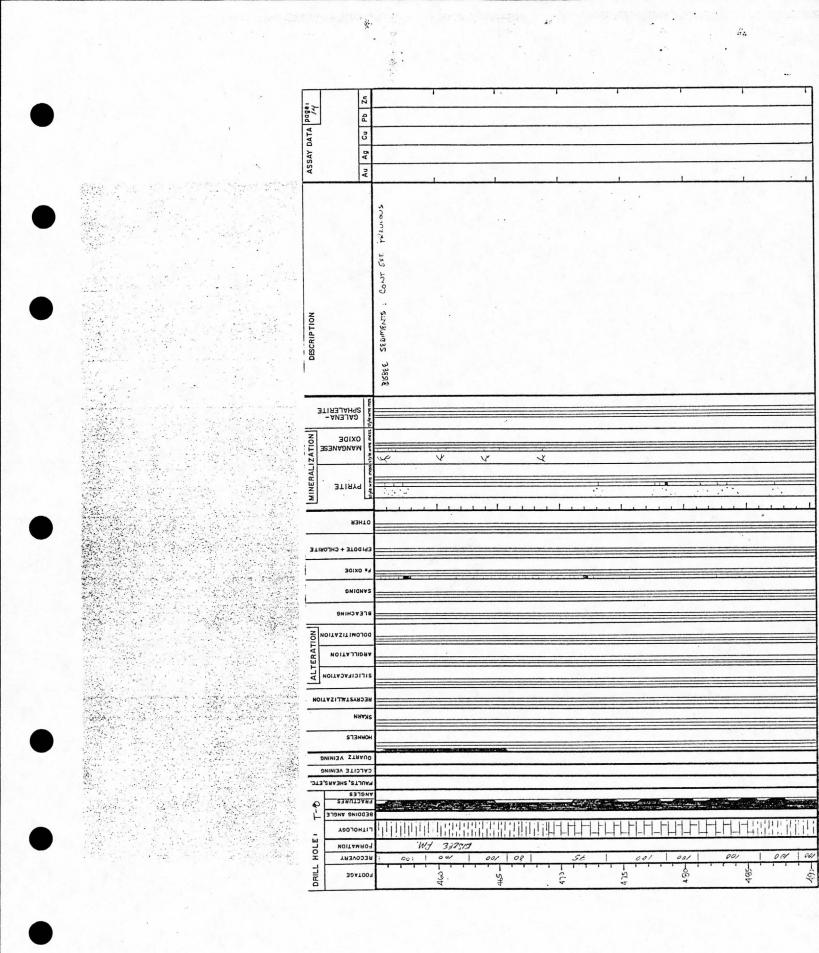
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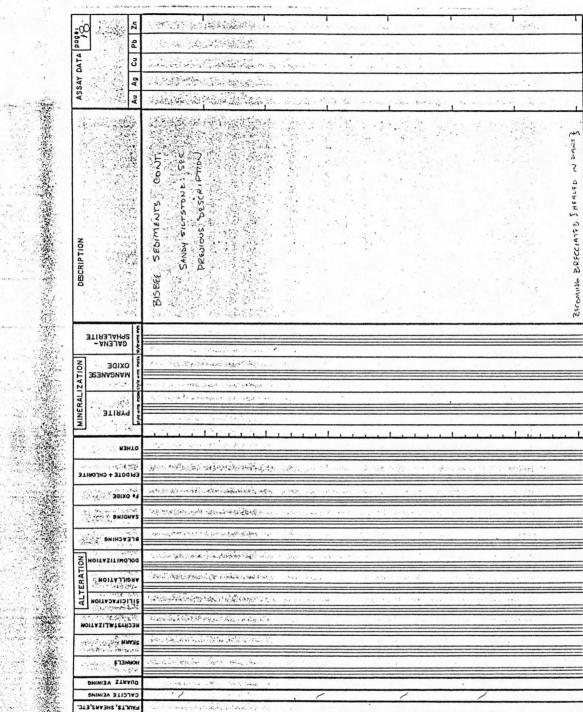
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