

CONTACT INFORMATION Mining Records Curator Arizona Geological Survey 416 W. Congress St., Suite 100 Tucson, Arizona 85701 520-770-3500 http://www.azgs.az.gov inquiries@azgs.az.gov

The following file is part of the

Richard Mieritz Mining Collection

# **ACCESS STATEMENT**

These digitized collections are accessible for purposes of education and research. We have indicated what we know about copyright and rights of privacy, publicity, or trademark. Due to the nature of archival collections, we are not always able to identify this information. We are eager to hear from any rights owners, so that we may obtain accurate information. Upon request, we will remove material from public view while we address a rights issue.

# **CONSTRAINTS STATEMENT**

The Arizona Geological Survey does not claim to control all rights for all materials in its collection. These rights include, but are not limited to: copyright, privacy rights, and cultural protection rights. The User hereby assumes all responsibility for obtaining any rights to use the material in excess of "fair use."

The Survey makes no intellectual property claims to the products created by individual authors in the manuscript collections, except when the author deeded those rights to the Survey or when those authors were employed by the State of Arizona and created intellectual products as a function of their official duties. The Survey does maintain property rights to the physical and digital representations of the works.

# **QUALITY STATEMENT**

The Arizona Geological Survey is not responsible for the accuracy of the records, information, or opinions that may be contained in the files. The Survey collects, catalogs, and archives data on mineral properties regardless of its views of the veracity or accuracy of those data.

# BOTTOMING OF THE UNITED VERDE SULPHIDE PIPE

 $\frac{By}{F}$ . Yates

PHELPS DODGE CORPORATION UNITED VERDE BRANCH JEROME, ARIZONA

> ARIZONA SECTION A. I. M. E.

TUCSON - OCTOBER 2(-28, 1946

# PHELPS DODGE CORPORATION UNITED VERDE BRANCH MINES DIVI 8ION

## "BOTTOMING OF THE UNITED V&RDE SULPHIDE PIPE"

By Paul F. Yates

#### INTRODUCTION

Ninety-five percent of the copper production of the United States has come from sixteen districts, of which J erome ranks sixth. "The United Verde sulphide pipe has been responsible for giving the Jerome district this rank. The United Verde Extension production was from a down-faulted segment of the upper part of the sulphide pipe. Production from other sources in the district is comparatively negligible.

It is estimated that the original sulphide mass had a total of more than 130 million tons and a vertical extent of over 8300 feet. The United Verde Extension segment accounts for approximately 1120 feet, and the United Verde, to the 4800 ft. level, 4770 feet. The balance was lost by erosion before and after the displacement of the United Verde Extension portion of the sulphide pipo.

Of the twenty most productive western metal mining districts, Billingsley and Locke in 1938,  $#$  listed eleven as having been bottomed, or thoroughly explored to the maximum justifiable depth. Four, including the Jerome district, were listed as not having been bottomed. The Jerome district must now be deleted from this latter category.

Production from the United Verde Extension ceased in 1937. Present production from the United Verde is largely from cleanup operations. Extensive exploration to a depth of 1500 feet below the lowest stoping lovel in the United Verde mine has shown this great sulphide. pipe diminishing into roots that are almost negligible.

The Jerome mining district is 1. Joated on the northeasterly 10pe of the Black Hills in central Arizona with the two principal mines, the United Verde and the United Verde Extension, close to the town of Jerome ncar the north end of tho district. The geology of tho district has been described by several writers. .\* Here it will suffice to review the more salient features pertaining to the environment of the United Verde sulphide pipe.

\*Copper Mining in North America, U.S. B.W. Bull.  $405$ , p. 2-8. U.S.B.M. Minerals Yearbook, 1944.

 $^{th}$ P. Billingsley and A. Locke, "Structures of Ore Doposits in the Continental Framework". A.I.M.E. Trans. 144, p. 9-64, 1941.

"Sce list of references.

 $\sqrt{1}$ (/

 $\frac{1}{2}$ 

 $\bigcirc$ 

#### GEOLOGY

The geological record begins well back in pre-Cambrian time with a vast outpouring of volcanic material of widely varying composition. This has been termed the greenstone complex and includes much rhyolitic flow and some fragmental material. Overlying this is the bedded sediments formation in which bedded volcanic tuffs and sedimentary material predominates. During a period of deformation the material was consolidated. folded, faulted and titled steeply to the northwest.

Rhyolitic (Cleopatra) quartz porphyry was then intruded as a large, somewhat tabular mass approximately along the greenstone-bedded The United Verde diorite followed in the form of sediments contact. a tabular blunt-nosed expanding plug roughly on the quartz porphyrybedded sediments contact. Often a remnant of the sediments was left between the porphyry and diorite. Locally the diorite cuts across porphyry tongues.

A series of mineralizing solutions followed the intrusion of the diorite, coming up through the marginal zone of the quartz porphyry. These solutions were controlled by the concave configurations of the overhanging diorite and the schistosity and fracturing in the quartz porphyry. The United Verde sulphide pipe was formed by replacement of part of the remnants of the sediments and much of the quartz porphyry by sulphides, quartz and ferruginous chlorite. After the formation of the sulphide pipe the region was subject to dynamic forces which resulted in major faulting. One of these faults, the Verde, which strikes northwesterly and dips to the east, resulted in the placement of the United Verde Extension orebody. It is estimated that the vertical displacement on this fault in prc-Cambrian time was something like 2400 feet.

Following a long period of crosion and peneplanation, the middle Cambrian (Tapeats) sandstone was deposited as a thin blanket (0-100 ft.) of ferruginous beach sand and pebbles which tended to fill the minor irregularities of the pre-Cambrian surface. Overlying the basal sandstone in the Jerome area are from 300 to 500 feet of Devonian limestone, 300 to 500 feet of Mississippian (Redwall) limestone, and from nothing to 500 feet of red (Supai) sandstone and shale of Permian age. The deposition of each of these periods was proceded and followed by periods of uplift and prosion. Later an outpouring of Tertiary lavas (Malpais) formed a covering mantle. Damming of the Verde Valley by the lavas and contemperancous subsidence permitted the deposition of over 2000 feet of impure while calcareous scidments (Verde formation).

Pre-lava crosion was evidently very active. A deep stream channel cut through the Paleozoic formations and partly exposed the gossan of the United Verde Extension segment of the sulphide pipe. This channel is now filled with Tertiary gravel and lava.

After the outpouring of the Tertiary lavas there occurred a period of normal faulting which relatively uplifted the west side of the Vorde Valley in the neighborhood of 5000 feet. The added displacement on the Verde Fault as indicated by the relative positions of the base of the Paleozoic formations in the vicinity of the sulphide pipe measures

approximately 1600 feet. The upthrust block containing the lower, or United Verde mine portion of the sulphide pipe, is bounded by four normal faults. The Verde to the east, the Warrior to the west, the Haynes to the north, and the Hull to the south. Post pre-Cambrian displacement on the Haynes Fault was 500 feet, the Warrior Fault 200 feet, and a comparatively small amount on the Hull Fault. In addition to the pre-Cambrian movement on the Verde Fault some evidence exists that there was also pre-Cambrian movement along the other three faults.

The uplifted scarps of the principal faults were especially subject to erosion with the resultant exposure of a strip of pre-Cambrian rocks paralleling the Verde Fault. The proximity of the exposed ore zone to the fault 'scarp induced rapid erosion which resulted in a comparatively shallow gossan zone above the United Verde mine. The secondary copper found. in the fault zone and near the surface to the east was evidently transported from the top of the sulphide pipe during this pariod.

#### SULPHIDE PIPE

#### Localization

The United Verde mineralized zone is located in the vicinity of a change in the original trend of the greenstone-bedded c diments contact. (Sec page  $4$ ). To the south from the mineralized zone the contact is in a nearly north-south direction and is comparatively regular. To the north the contact swings to the northoast and is very irregular and interfingering, also to the north near the ore zone, the bedded sediments show many small drag folds evidently related to the plane of weakness along the greenstone-bedded seciimcnts contact. This crumpling is more intense in the upper part of the mine than in tho daeper levels. It is believed that the major structure is a large drag fold which plunges steeply to the north-northwest, is more open with depth, and becomes progressively tighter and stronger with increased elevation. The attitude of the minor drag folds support this hypothesis. Much of the evldence along this contact has been destroyed by the intrusion of the quartz porphyry and diorite. The form of these intrusives was affected by the steeply pitching folds, particularly in the. case of the quartz porphyry.

The porphyry shows a variable schistosity or parting that is bolicved to be the result of differential pressure before the rock was completely solidified. The relatively local zones of more intense schistosity were formed later. To the south and cast the schistosity strikes west of north and dips to the east. Near the mineralized zone it strikes in a broedly curving arc tending to parallel the surface of the over-riding diorite. The variable schistosity, with zones of greater shearing permitting greater penetration of mineral solutions, is responsible for the irregular footwall of the mineralized zone. In general the amount of shearing near the minoralized zone incroasos with clevation.

The diorite came in pushing upward and to the northeast, and when it had expended sufficiently, wrapped around the bend in the porphyry contact. On the bottom levals the east diorite contact is in a north-south line with a bulge or flare to the east near the northern part. Below the



3000 ft. level the contact dips to the east to parallel the schistosity of the porphyry. Above the 3000 ft. level, as the diorite expands, the dip is to the west or across the schistosity of the porphyry. As higher elevations are reached, the bulge to the east becomes progressively more prominent. In the middle levels it forms a gently curving arc, and in the upper part of the mine a point swings to the south to form an inverted structural trough. (See plans and vertical section). The plunge of the trough is northerly conforming to the general plunge or the axes of the minor drag folds in the sediments.

#### Minor Faults and Breaks

The quartz porphyry near its margin with the diorite is cut by an irregular branching pattern of faults. In the upper part of the mine they are either weaker or obscured by mineralization and are hard to trace. In the lower part of the mine they can be traced over considerable distances. These faults appear to have acted as a partial control for the mineralizing solutions. It is significant that the sulphide masses are on the hangingwall side of the breaks, or between the breaks and the diorite. The solutions which formed the black schist penetrated to a certain extent into the footwall side of the breaks, though cross breaks or fracturing associated with the cross breaks. These faults are presumed to have been formed by the forces associated with the intrusion of the expanding diorite as it pushed or wrapped around the bend in the porphyry contact.

The principal breaks of this fault pattern strike northeasterly and are found starting near the diorite to the south and diverging from it as they skirt the footwall of the sulphide with some branches swinging more to the east and others back toward the diorite. They dip to the northwest at a flatter angle than the diorite-porphyry contact and thus as higher elevations are reached, the space between the breaks and the diorite increases, but still dip toward the contact. (See vertical section Page 6).

A second, less important series, strike in a northwesterly direction and dip to the northeast. In several places the breaks of this series appear to be responsible for the localization of mineralization in the quartz porphyry footwall. Breaks of this series also conform to the southwesterly boundary of the northeast portion of the main mineralized zone.

Faults of both series throw out complicated spurs and branches. The breaks vary in width from a few inches to as much as two or three feet, with crushed rock, quartz, carbonate and gouge filling. They are locally, if not usually, mineralized with pyrite ranging from scattered crystals to streaks and patches. In a number of cases dikes have followed these breaks. The faults are definitely carlier than the intrusion of the dikes and the later mineralization and are believed to be entirely pre-mineral. It is probable that the forces which produced these faults, fractured the already schistose porphyry located between the breaks and the diorite thus forming a more favorable host for the mineralizing solutions.

#### Dikes

Numerous nearly vertical dioritic dikes cut through the orezone and surrounding rocks in a general east-west direction. In width they gonorally vary from a fraction of an inch to about three feet. However, there are a number six to eight feet wide, with exceptional cases of

Pago 6.



 $\sum_{i=1}^{n}$ 

 $\bigcap$ 

greater width. In some places they are quite regular, in others irregular and broken. The increase in number and local bunching of the dikes, near copper bearing areas, is noticeable with depth. A series of larger dikes starting near the 1500 ft. level cut through the sulphide mass roughly paralleling the diorite contact. With depth they pass towards the footvvall of the mineralized zone and on the 3300 ft. level are almost entirely in the chlorite schist and porphyry.

The smaller dikes are usually very fine grained and in places altered or partially altered. The larger dikes are usually fresh and of a coarser texture. Most of the dikes appear to have been intruded following the principal stage of copper mineralization, and before the later stages. There is a possibility that some of the dikes in the lower part of the mine are older.

## Mineralization ,

.. ,.

'J

The United Verde mineralized zone consists of a very irregular pipelike body of massive sulphide, quartz and mixed sulphide and rock with a steep north-northwesterly plunge. It is clearly of the schist replacement type, with replacement of schistose quartz porphyry and a part of the fringe of the bedded sediments. Replacement has been so complete that the relative amounts of the two rocks cannot be determined exactly, but the available evidence indicates that the porphyry was in considerable excess. The cross s0ction of the mineralized zona varies throughout the mine. From the surface to the 1200 ft. level, the horizontal cross section averages about 250,000 square fect. From this horizon to the 1650 ft. level it averages about 500,000 square fect. From the 1800 to the 3000 ft. level it averagos a little over 400,000 square foot. From the 3150 ft. lovel down, the total area decreases rapidly and at the 4500 ft. level it only totals 37,000 square feet, of which 9700 square feet is massivo sulphide.

The sulphide zone is limited to the north and  $m$ st by the impervious diorite, and to the south and east by an irregular border of black chlorite schist. The most abundant primary gangue mineral is pyrite which makes up the great bulk of the sulphide mass. With it are associated large quantities of quartz, chlorite and dolomite. Chalcopyrite is by far the principal copper bearing mineral. Less important associated minerals are sphalerite (marmatite), tennantite, bornite, galena and specularite. The minerals of the oxide zone and zone of a secondary enrichment have been included in earlier reports and aro not repeated horo.

Several stages of mineralization arc recognized and are briefly described as follows<sup>\*</sup>:

1. The first solutions followed the relatively permiable shear zones in the porphyry adjacent to the overhanging diorite, replacing much of tho intervening porphyry and portions of the fringe of bedded sediments. Those solutions doposited large quantitics of quartz with very minor quantities of pyrite and chalcopyrite.

\*I. E. Reber, "Jerome District", Some Arizona Copper Deposits, Arizona Bureau of Mines Bull. 145, Pages 41-65, November, 1938.

I..

2. Solutions of the second period deposited much pyrite with important quantities of marmatite, a little chalcopyrite and probably local quartz and dolomite. This period is responsible for the major part of the sulphide zone and probably much of the zinc.

3. In the third period solutions working out from the footwall of  $\vee$ the pyrite replaced large quantities of porphyry with a nearly black, high iron variety of chlorite.

4. The fourth period of mineralization was responsible for most of the commercial ore and consisted largely of chalcopyrite with minor pyrite and a little intergrown marmatite and galena. The chalcopyrite appears to have favored the replacement of the chlorite schist.

5. Solutions of the fifth period brought in small quantities of quartz, in part associated with bornite and probably other sulphides.

6. The sixth period of mineralization was characterized by the deposition of intergrown quartz and dolomite with associated chalcopyrite, pyrite, tennantite and sphalerite. These minerals are normally found in the chlorite schist near its footwall contact with the porphyry and most abundantly in the lower stoping levels. In this period mineral deposition by replacement was followed by deposition in fractures and gash veinlets. This mineralization has added materially to the schist ore in certain localities.

7. The latest phase of mineralization resulted in widespread deposition of quartz and dolomite with sparse pyrite in small veins and gash veinlets.

The products of each succeeding phase of the mineralization cycle built up on the footwall of the previous phase, and while rock replacement was most important, some of the products of the preceding phase were nearly always replaced.

In the lower levels of the mine the diorite and bedded sediments parallel the schistosity of the quartz porphyry and the ore zone consists of a few disconnected sulphide-schist lenses. These lenses are in the porphyry near its margin. Two of these are most significant. The lonso north of the bulge in the diorite is most extensive on the 4500 ft. level. It is a sulphide-schist lonse 280 feet long by about 45 feet wide, with a tail stringing out to the north. Its cross section nearly doubles by the time it reaches the 4050 ft. level, continues about the same to the 3750 ft. level, and then decreases rapidly to pinch out under the expanding diorite just below the 3300 ft. level. The ore in this lense is limited to a cross section of 800 square feet on the 4500 ft. level, increases to 1300 square feet on the 3750 ft. level, then decreases rapidly above this horizon.

Immediately south of the bulge in the diorite on the 4500 ft. level, opposite a rather sharp embayment in the contact, is a lense of quartz, sulphido and schist, 150 feet long by about 30 feet wide. This lense rakes sharply upward to the south, increases rapidly in size to the 3750 ft. level, continues about the same size to the 3450 ft. level. It then decreases up to the 3300 ft. level with a part continuing above the 2700 ft. level. There is very little copper in the massive sulphides in this lense. The copper ore

is in the chlorite schist localized along one of the branching faults and separated from the massive sulphides by a band of quartz porphyry, or altered material, transitional between porphyry and chlorite schist. The principal ore body in this lense extends from the 3300 to the 2700 ft. level. There are other mineable bodies of limited size between the 3450 and the 3150 ft. levels.

There is a small low grade irregular pipe-like lense of quartz, massive sulphide and mineralized schist associated with quartz-porphyry in a sharp embayment on the west side of the diorite. The material differs only from that of the main ore zone in the presence of visible magnetite and pyrrhotite. This lense pinches out against the overhanging diorite above the 2700 ft. level. It plunges easterly toward the main ore zone, suggesting that it is a small branch comparable to the "north lense".

The intervening diorite makes the upper part of the north lense a distinct branch or spur of the mineralized zone, while the south lense may be looked upon as an offset portion of the main ore zone with a vertical overlap of more than 1000 feet on the main sulphide pipe.

.. '

 $\bigcirc$ 

On the 4500 ft. lovel, scattered schist areas with minor sulphide, underlie the main sulphide mass. They extend upward, become stronger and make the bottom of the deopest clement of the main sulphide mass, which comes in some distance below the 3750 ft. level. The east clement extends from below the 3300 ft. level to join the main mass near the 3150 ft. level; the south element, from below the 3150 ft. level to join the mass near the 2850 ft. level. From the 2850 ft. level upward, the sulphide pipe is continuous.

The start and upward expansion of the east element of the main sulphide body is related to a band of increased schistosity in the quartz porphyry, which grows progressively stronger upward. This zone of schistosity, when combined with the control effected by the expanding bulge in the diorite, accounts for the formation of the crescent shaped sulphide mass in the middle levels. As the diorite continues to expand and over-ride the porphyry, the mineralizing solutions are progrossively more confined until in the upper levcls, the sulphide is nearly cliptical in form, with the major axis almost perpendicular to tho crescent.

Copper. Since the copper mineralization appeared near the close of the mineralizing cycle, most of the ore occurs on the underside or footwall of the sulphide mass. In the lower horizons there are only minor spots. The copper mineralization becomes more effective as the higher elevations in the mine are reached, and on the 3300 ft. level, there are three mineable bodies and the bottom of a fourth. In the lower stoping levels most of the copper is found in the chlorite schist. In the upper levels the amount of copper in the schist decreases and progressively more is found in the massive sulphide. On the 2850 ft. level the copper mineralization was more widespread than below, but the larger stopes are in the chlorite schist near the split in the sulphides. Above this horizon the ore bodies become progressively more continuous and on the 2100 ft. lovel the ore bend is 15 to 90 feet wide and 800 feet long. The flattening of the overhanging diorite permitted increased mineralization near the 1800 and 1650 ft. levels. In the uppor horizons of the mine the heavy pyrito stage of mineralization left unreplaced stringers of rock extending far out into the sulphide. Thus with tho more constricted diorite trough and a more intense zone of shearing in the porphyry, the heavy chalcopyrite mineralization penetrated nearly to the diorite hangingwall, as well as extending far out into the quartz porphyry footwall.



 $\gamma^{\prime}$  $\tau_{\rm c}$  $\mathcal{A}_\mathcal{I}$ ∵'−  $\mathcal{L}$ 小 ÷  $\mathcal{P}^{\mathcal{C}}$  $\langle \rangle$  $\frac{1}{2}$  $\mathcal{L}_{\mathcal{F}}$  $\tau^{\Lambda}_{\mu\tau}$  $\tau^{\prime}\tau$  $\mathbb{C}^{\mathbb{R}}$  $\frac{1}{2}$ रंग **SEDIMENTS**  $\frac{A}{\sigma_{\rm s}} \leq$  $\sigma^2$  $\frac{1}{2}$  $\mathcal{L}_{\mathcal{L}}$  $\mathbb{R}^2$  $\mathcal{N}_\mathrm{c}$  $\frac{1}{2\sqrt{2}}$  $\ddot{z}$ 43 EDIMEN T S S  $\mathcal{L}$  $\mathbf{z}^{\mathbf{t}}$  $\ddot{\phantom{a}}$  $\mathcal{L}$ ⇔  $\sigma_{\rm eff}^{\rm tot}$  $\mathcal{M}$  $\div$  $\mathcal{L}$  $\mathbf R$  $\mathbf{I}$ Е D  $\Omega$ Т τ  $\mathbf I$  ${\bf R}$ I D o  $\sigma_{\rm eff}^{\rm 20}$  $\frac{1}{2}$ λF  $\lambda_{\rm S}$  $\mathcal{L}$  $\sim$  $\mathcal{L}$  $\mathbb{R}$ مبيء SULPHIDE  $\tau^{\prime}$  $\zeta_{\rm D}$  $\gamma_{\rm s}$ **SULPHIDE**  $\ddot{\sim}$  $\mathcal{L}$  $\mathbf{x}$  $\mathbf{x}$  $\bar{\mathbf{x}}$  $\frac{A}{12}$ Scr ÷  $\ddot{\mathbf{v}}$  $\mathcal{L}$  $\mathcal{L}_{\mathcal{L}}$  $\mathbf{x}$  $\mathbf{x}$  $\mathcal{P}_\mathcal{L}$ QUARTZ  $Q$ U A R T Z  $\sum_{i=1}^n \mathcal{E}_i$  $\mathbf{v}^*$ PORPHYRY ×  $\dot{\texttt{P}}$ ORPHYRY ÷.  $\mathbf{v}_k$  $\mathbf{x}$ Pago ∾  $2850$ <br> $1" = 330'$  $2100$  $\mathbf{x}$  $\mathbf{x}$  $2<sup>n</sup> - 330<sup>n</sup>$ 



# CHART OF MINERALIZED AREAS

## Thousands of Square Feet

₽,



Pago 13.

Zinc. As already stated most of the zinc is associated with the principal stage of pyrite deposition before the significant copper deposition. It is found in the massive sulphide between the zone of copper deposition and the diorite hangingwall. The greater concentration of zinc is found in the upper levels with scattered spots down to the 3000 ft. level. Below this, relatively high zinc values are limited to part of the north sulphide lense between the 3750 and the 4500 ft. levels.

Development of the low grade zinc ore disclosed that it is quite irregular and discontinuous. Much of the best of the zinc has either been removed in the open pit operation or is involved in the general mine subsidence. The intimate mineralogical composition indicates a difficult metallurgical separation. These factors combine to preclude any present possibility of commercial zinc production from the United Verde mine.

D,

Precious Metals. The best primary gold and silver values are associated with relatively siliceous material. A number of the massive sulphide copper ore bodies terminate in siliceous material, with relatively high gold and silver values. There are also localized gold-silver values in the comparatively narrow transition zone between lean sulphide and quartz. The average ores carry in the neighborhood of .015 ounces of gold per ton and a little over an ounce of silver. Increased gold may go to as much as .10 ounces and silver to two or three ounces. Occasionally much higher values have been found. Silver averaging about two ounces is found in schist ores with comparatively abundant tennantite.

Chlorite Schist. In the upper and middle levels the preponderance of the chlorite schist is nearly pure ferruginous chlorite and almost black in color. In the lower stoping levels and progressively below, the amount of material that is transitional between the quartz porphyry and the chlorite schist continues to increase. There are sizeable areas where it is nearly impossible to tell whether chlorite schist or porphyry predominates. It is probable that in the lower horizons the paths of the mineralizing solutions were more scattered and hence much of the porphyry was only partially In the middle and upper levels, structural conditions led to a replaced. greater concentration or localization. There is a progressively marked lessening of the amount of chlorite schist developed on each of the succeding levels from the 3300 ft. level down.

Quartz. The large masses of quartz found underlying the diorite hangingwall are normally quite dense and commonly jaspery in texture. It is suggested that this earlier quartz assisted in rendering the hangingwall impervious. Other quartz masses are found throughout the sulphide. In general, as the diorite contact is approached, the amount of quartz increases. On the bottom development levels, there is a notable increase in the amount of quartz as compared to the heavy sulphide. Jaspery quartz lenses, varying from a fraction of an inch to several feet in thickness, are often found in the quartz porphyry underlying the sulphide masses and in shear zones adjacent to some of the minor faults. These lenses are usually mineralized with scattered pyrite and a very little chalcopyrite.

### ORIGIN AND CONTROL OF MINERALIZATION

General knowledge of metalliferous ore deposits makes it practically certain that the United Verde sulphide pipe was formed by solutions related to igneous rocks or an igneous magma reservoir, though the particular magnatic source to which the mineralizing solutions owes their origin is speculative. The location of the Jerome district with respect to the Bradshaw granite area to the south and southwest, and the granitic areas to the west and northwest, makes the Bradshaw granite magma the most plausible source of the mineralizing solutions."

It appears that the mineralizing solutions were introduced through the same channel or zone along which the quartz porphyry and diorite were intruded. As the north end of the diorite continues to recede to the south with depth, it exerts less and less control on the deposition of the solutions. There is every indication that at some depth below the 4500 ft. level the mineralizing solutions were introduced along a nearly north-south striking shear or break that is on or close to the bedded sediments-quartz porphyry contact.

It is believed that the breaks of the minor fault pattern related to the intrusion of the diorite, were most significant in the control of the mineralizing solutions between the postulated major break at depth and the zone of effective mineral deposition. Various unfavorable factors, such as the weakness or fluidity of solutions, tightness of channels, and lack of the damming effect of the overhanging diorite, prevented more than a limited amount of deposition in the bottom levels of the mine. With increased elevation the more sheared and fractured porphyry, coupled with the changing composition of the solutions, and the increased damming effect of the diorite permitted more and more material to be deposited. The diorite, the fringe of sediments and the earlier quartz, combined to form a dam that was impervious to the later phases of the mineral solutions. The changes in the contour of the diorite, the variable degree of schistosity in the porphyry, and the habit of the breaks all combined to effect changes in the form of the mineralized zone so that no two levels are the same.

### DEEP LEVEL EXPLORATION

When underground operations were resumed in 1937 as the open pit was being completed, the lowest stoping was on the 2550 and 2700 ft. levels, the 2850 and 3000 ft. levels were only partially developed, and drifting had just been started on the 3300 ft. level. The necessity for additional exploration was apparent if the proposed mining schedules were to be maintained.

Preliminary drifting on the 3000 ft. level was completed, and the heading on the 3300 ft. level was being continued along the sulphide-schist contact. An extensive diamond drilling program on the 000 ft. level indicated a decrease in the area of mineable ore as compared to the levels immediately above. As the 3300 drifting progressed, attendant drilling soon indicated an apparent breaking up of the sulphide mass, with conspicuously little copper mineralization. This suggested the possible termination or bottoming of the sulphide pipe.

\*L. E. Reber, op. cit.

 $Page116$ .  $N_{\times}$  $\mathbf{x}^{\top}$ 夸 أتأثر  $\mathcal{A}_{\mathcal{P}}$ SEDIMENTS سنوا  $\mathcal{E}$  $\sim$  at the de ar 第二卷 要 第二章 表示  $\mathcal{L}^{\text{L}}_{\text{L}}$ 선  $\mathbf{E}$  $\overline{\mathbf{x}}$  $\langle \cdot \rangle^{\rm L}_{\rm L}$ 第二卷 花 表示 多、多、分  $\gamma_{\rm c}$  $x = 0$  ,  $y = 1$  $\mathbb{Z}$ 1. 要 一定 QUARTZ 一、音、一、音。 K PORPHYRY 第二号 2 经收益 × 第二章 第二卷 经 第二卷 作者 會 要 第二卷 第三卷 卷  $\sim 25$  .  $\label{eq:3.1} \frac{d\mathbf{r}}{d\mathbf{r}} = \frac{d\mathbf{r}}{d\mathbf{r}} = \frac{d\mathbf{r}}{d\mathbf{r}} = \frac{d\mathbf{r}}{d\mathbf{r}} = \frac{d\mathbf{r}}{d\mathbf{r}}$  $\mathcal{L}^{\text{L}}_{\text{c}}$  $\mathcal{A}_{\mathcal{I}}$  $\mathcal{L}^{\text{max}}_{\text{max}}$  $\mathcal{L}_{\mathcal{C}}^{(k)}$  . 一卷  $4\,$  5  $1'' = 400'$ 

 $\mathcal{L}_{\mathcal{L}}$ SEDIMENTS.  $\hat{\mathcal{C}}$  $\frac{1}{4}$  $\mathcal{L}$  $\mathcal{L}_{\mathcal{L}}$ ×.  $\pmb{\times}$  $\frac{1}{4}$  $\mathcal{A}^{\text{L}}$  .  $\Rightarrow$  $\frac{d^2\mathbf{r}}{d\mathbf{r}}$  .  $\sqrt{\dot{x}}$  $\tilde{\mathbf{X}}$  $\frac{1}{2} \frac{1}{2}$  .  $\mathcal{A}_{\mathcal{I}}$  $\frac{1}{2} \mathbb{Z}$  $\overline{\mathbf{x}}$  $\gamma^{\rm t}$  :  $\vec{r}$  $\sigma_{\rm c}^{\rm b}$  $\mathcal{L}^{\mathcal{L}}_{\mathcal{L}}$ 는  $\mathcal{D}$  $\ddot{\cdot}$  $\sigma_{\rm c}^{\rm MS}$  $\psi_2$ Ń,  $\langle \hat{r}_s^{\prime}\rangle$  $\gamma^{\rm L}_{\rm L}$  $\gamma^{\rm th}$  $\mathcal{L}$  $D$   $I$  $\bullet$  $\mathbf R$  $\mathbf T$  $\mathbf T$  $\mathbf E$  $\bar{\mathbf{x}}$ **カートホートボートボ**  $\overline{\mathbf{x}}$ 第二章 表  $\bar{\mathbf{x}}$  $\bar{\mathbf{x}}$  $\overline{\mathbf{x}}$ 在 一步  $\mathcal{A}_{\mathcal{G}}^{\mathcal{G}}$  , where  $\mathcal{G}_{\mathcal{G}}$ 것이  $\mathbb{Z}^{\mathsf{tr}}$  $\pm$  $\mathbf{r}$ QUARTZ (我) 第二种 一花 PORPHYRY £  $\mathbf{x}$ x 水。 水。 水。 水。 Ü ا<br>منابع  $\bar{\mathbf{x}}$  $\bar{\mathsf{x}}$  $\pmb{\times}$ 教育 教育学校  $\mathcal{L}_{\mathcal{C}}$ 救 我 经  $\label{eq:2} \mathcal{L} = \mathcal{L} \mathcal{L} \mathcal{L} \mathcal{L} \mathcal{L} \mathcal{L}$  $\mathbf{x}$  $\prec$ 秦、 我 一 我 。  $\sigma_{\rm m}^{\rm M}$  and  $\sigma_{\rm m}$  $\bar{x}$ ×, 第二卷 & 文  $\mathcal{L}$ **2. 经工厂经营** ×.  $\mathbf{x}$  $\bar{\mathbf{x}}$ 750 3  $1'' = 400'$ 

ago 18 SEDIMENTS  $\mathcal{L}(\mathcal{L})$  $\mathcal{L}^{\mathbf{c}}_{\mathbf{a}}$  .  $\mathcal{L}_{\mathcal{C}}$ Y.  $\mathbf{\hat{x}}$ <u> 동</u>화 1979년 - 대한민국의 1979년 1월 1989년 1월 198 - 六一  $\frac{1}{2}$  $\mathbf{x}$ 一、蜂  $\frac{1}{2}$ 最高 率 **处** 一致 一、数  $\mathcal{L}^{\text{max}}_{\text{max}}$  $\mathcal{L}^{\prime}$  $\mathbb{Z}_\ell^{\bullet}$  :  $\mathcal{L}^{\mathcal{L}}$  . D I O R I T E 第二条 1 **蒙**  $\gamma^{\prime}$  :  $\mathcal{F}_{\mathcal{G}}$ 一本 一章 一章 x × 1. 保证 一张  $\frac{d}{dt} \phi$ 2. 女 2. 一女 2. 永 家。来。  $\sigma_{\rm eff}^{\rm M}$  (  $\sigma_{\rm eff}$  $\mathcal{L}^{(1)}_{\mathcal{M}}$ 一起 经  $\frac{1}{2}$ (1) 我 (1)  $\frac{1}{\sqrt{2}}$ 一定 之  $\mathsf{x}$  $\pmb{\times}$ × ×  $\frac{1}{\sqrt{2}}$  $\overline{\phantom{a}}$  $\overline{\mathcal{X}}$  $\overline{K}$  $\bar{\mathbf{x}}$  $\frac{1}{2} \frac{1}{2} \frac{1}{2}$ **一定** 数 一  $\frac{1}{2}$ QUARTZ<sup>x</sup> **PORPHYRT** 第二章 经营业  $\bar{\mathbf{x}}$  $\mathbf{x}$  $\pmb{\times}$  $\dot{\mathbf{x}}$ 家。火。 安  $\mathcal{L}$  $\pmb{\times}$  $\pmb{\times}$ ×. **Participant** 一个  $\pi^0_{\rm{tot}}$  $\frac{1}{2}$ 物。 火。  $\mathcal{L}^{\text{L}}$ **1000 号**  $\mathbf{\hat{x}}$  $\mathbf x$ **如 故 那 的** 一款 一 经 一 六  $\mathcal{L}(\mathbf{r})$  $345$ ( )  $1'' = 400'$ 

An intensive exploration and development program was essential to delineate mineable ore areas below the 3000 ft. level and to determine whether or not the ore zone had definitely bottomed. A study of geological data indicated the split in the sulphide mass as the most favorable location for a development shaft. A new shaft was necessary because the existing shafts were at an excessive distance from the mineralized zone and because the hoisting equipment was inadequate for development at the desired depths. The new, No. 8 Shaft, was started from the 3000 ft. level in September 1939, and completed at the 4631 it. level in August 1942.

Exploration levels were established on the 3450, 3750 and 4500 ft. levels and later an intermediate on the 4050 ft. level. The extensive development and the areas prospected on the 3450, 3750 and 4500 ft. levels are shown on pages 16, 17, and 18. Development from January 1940, to October 1946, has included 10,230 feet of drifts, and nearly 17 miles of diamond drilling. The No. 8 Shaft exploratory program will be completed early in 1947 at a total cost of approximately \$1,000,000.

€

đ

ಕ

 $\frac{1}{2}$ 

The elevation of the 4500 ft. level is 1030 feet above sea level. Rock temperatures encountered as the development headings on this horizon wore being extended were 110° F.

#### **CONCLUSION**

It is believed that the great United Verde sulphide pipe has bottomed, that an extensive exploration program was entirely justified, and that it has been diligently carried out. Probably the one outstanding feature of the whole exploration program has been the succession of negative results ob-The minor spots of minoralized material encountered are a part of tained. the progressive breaking up and diminution of mineralization below the 2400 ft. level. The present exploration program has included a vastly extended area for a vertical distance of over 1500 feet below the lowest stoping horizon with essentially negative results. The chances are exceedingly remote for a change in structural conditions which would permit the deposition of mineral solutions at some deeper horizon.

Mining oper tions are now essentially confined to the cleaning up of ore in the vertical and horizontal pillars left by earlier mining, to remnants around the borders of existing stopes, and the completion of a few open stopes on the 3150 and 3300 ft. levels. The combined segments of the United Verde sulphide pipe, (United Verde and United Verde Extension), have produced  $34,550,000$  tons of ore from which has been extracted  $3,551,000,000$  pounds of copper, 53,420,000 ounces of silver, and 1,462,000 ounces of gold. This vast ore zone, with over sixty years of active productivity, can now be said to be effectively bottomed.

#### APPRECIATION

The writer is greatly indebted to Dr. Louis E. Reber, Chief Geologist of the Phelps Dodge Corporation, for his generous assistance in the editing of this paper. To Mr. J. B. Pullen, General Superintendent, and Mr. C. E. Mills, Mine Superintendent, of the United Verde Branch, Phelps Dodge Corporation, for their criticism and help; to Dr. C. A. Anderson, Geologist, United States Geological Survey, for his criticism. Acknowledgement is also due the officials of the Phelps Dodge Corporation for permitting the publication of this paper.

#### **REFERENCES**

i---------- ----------:--~--------,:\_\_---:-------------.

- I. E. Reber: "Geology and Ore Deposits of the Jerome District", A.I.M.E., Trans. LXVI, p. 3-26. (1922).
- Waldemar Lindgren: "Ore Deposits of the Jerome and Bradshaw Mountains Quadranges". U.S.G.S. Bull. 782. (1926).
- M. G. Hansen: "Geology and Ore Deposits of the United Verde Mine" Min.- Cong. Journal XVI, p.  $306-310$ .  $(1930)$ .
- L. E. Reber: "Jerome District", Some Arizona Copper Deposits, Ariz. Bur. Mines Bull. 145, p. 41-65. (1938).
- E. N. Pennebaker: "Review of Geology and Exploration Possibilities at the United Verde Mine", unpublished reports.  $(1943-45)$ .

J aroms, Arizona October 24, 1946

/

'0

f

 $Q_{\odot}$ 

đ

o

*A*  {j

 $\frac{1}{2}$ 

J •

*l*